

On the influence of coherent structures on flow hydrodynamics, transport and mixing at river confluences

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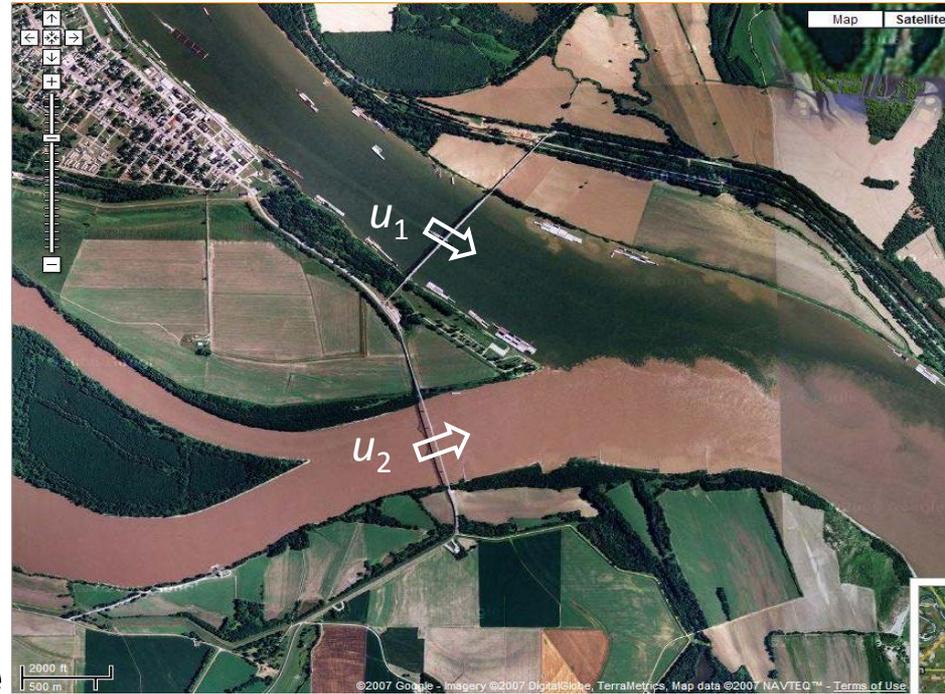
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Are fancy numerical simulations of any value for a river mechanics person?

INTRODUCTION

River confluences

- fundamental elements of natural drainage networks
- play an important role in regulating the movement of sediment through braided river systems
- are habitats of high ecological value



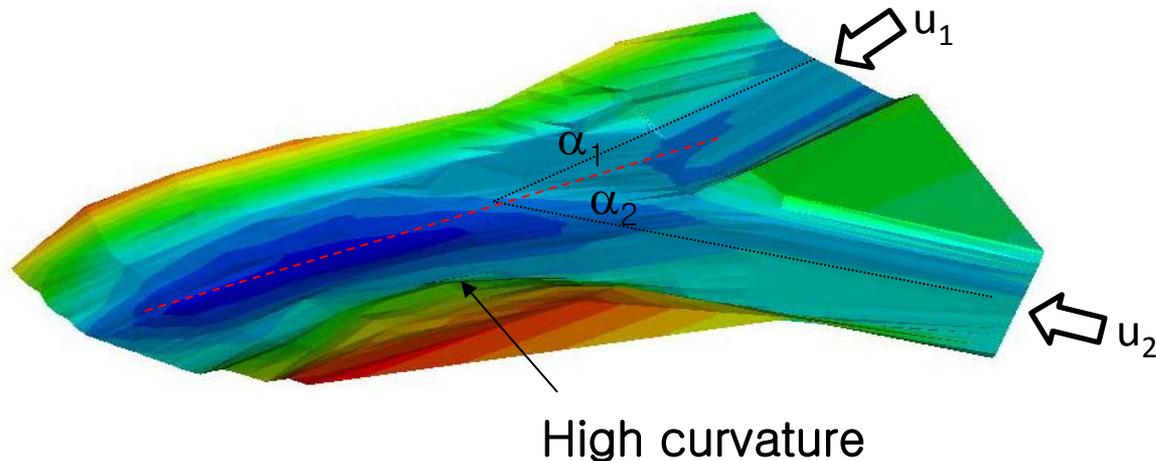
- dynamics of mixing controls how tributary inputs of nutrients and food are dispersed within a main river

General features:

- large-scale coherent structures form inside and around the MI
- flow conditions are fairly shallow

MAIN PARAMETERS

- **velocity ratio** $VR=U_1/U_2$
- **angles** between the two incoming streams and downstream channel
- degree of **concordance of the channel bed**



THE PLAN

-explore a system where large-scale coherent structures control mass exchange, mixing and transport processes

-try to validate/amend some hypotheses advanced based on limited amounts of field/laboratory data

-go from simple to complex:

**-sometimes we can learn a lot from less complex cases
(fewer variables, simpler to understand and parameterize)**

-be aware these less complex cases are not relevant for all types of natural stream confluences

MIXING INTERFACES AT RIVER CONFLUENCES

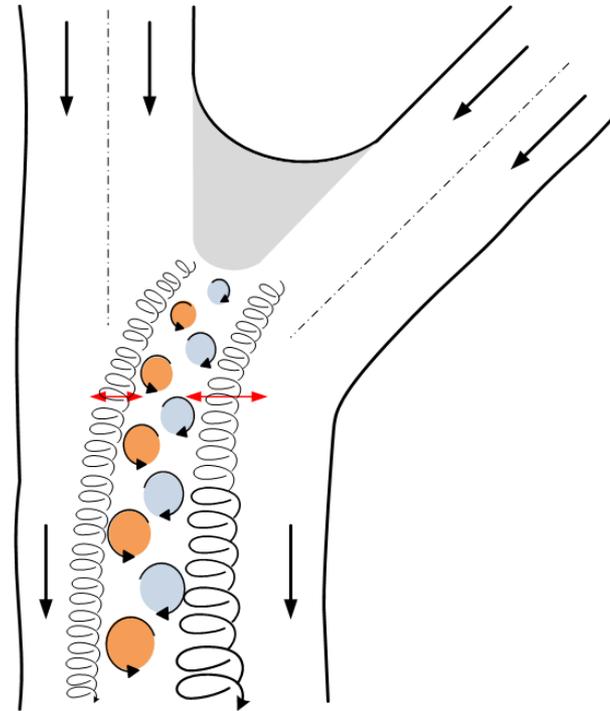
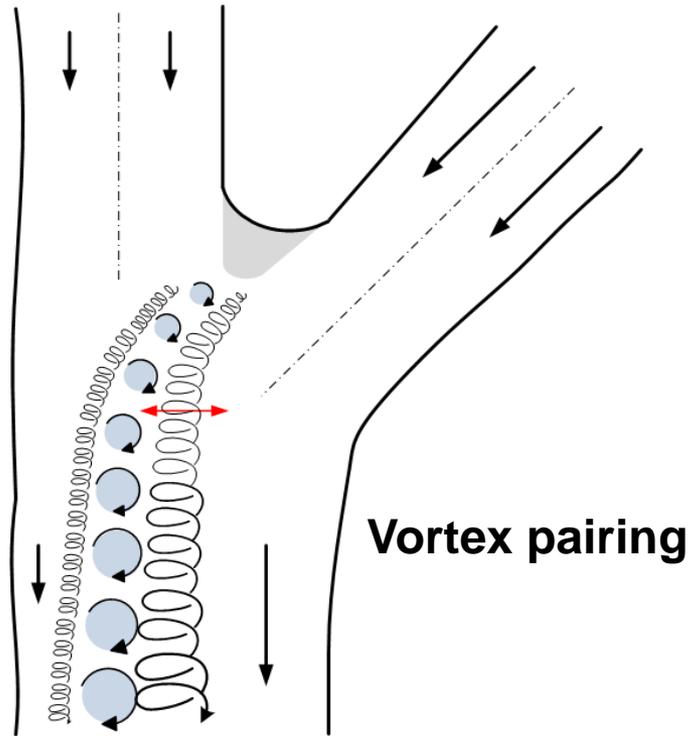
- Except maybe for confluences with a high degree of bed discordance, the region where the two streams come into contact can be described as a **shallow MI containing quasi-2D eddies**
- At river confluences, development of MI is strongly affected by **bed friction**

Classification of Mixing Interfaces

(Constantinescu et al, WRR 2011, JGR 2012)

Kelvin-Helmholtz mode ($VR \gg 1$)

Wake mode ($VR \sim 1$)



MIXING INTERFACES

-If the **angle** between the two streams is **small** and the **velocity difference** between the two streams is **high**, MI resembles a

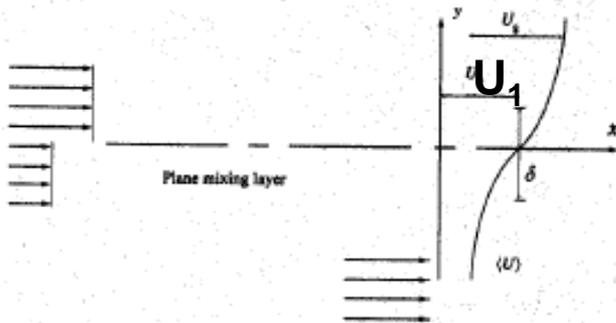
SHALLOW MIXING LAYER developing between two parallel streams

-This is the simplest type of MI developing at a river confluence

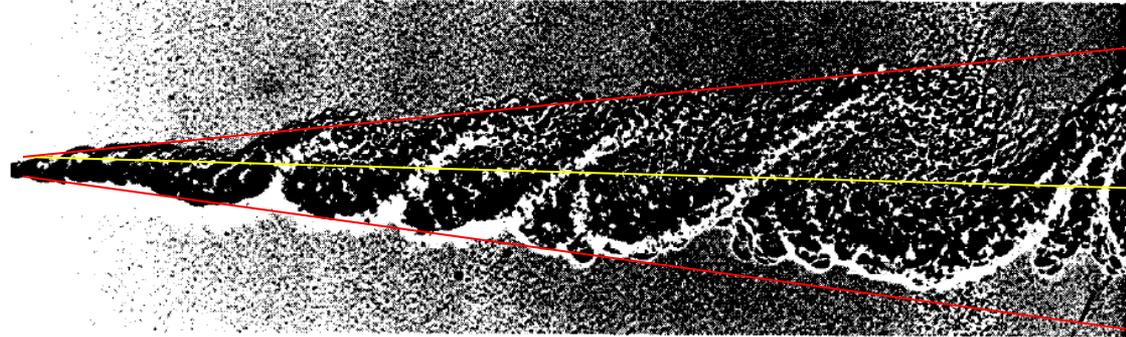
-Some fundamental results:

BACKGROUND

Simplest case of ML: 'Deep' ML



U_2



Visualization of structure of turbulent ML
(Brown and Roshko, JFM 1974)

-Mixing layer is self-similar

Spreading Rate $\frac{d\delta(x)}{dx} = const = \frac{\Delta U}{U_c} S'$ $S' \sim 0.09$

-A deep mixing layer grows linearly with no bound!

-Vortex pairing produces larger and larger KH billows

Confluences between **parallel** streams

Main parameter describing spatial development of a shallow ML:

S – bed friction number

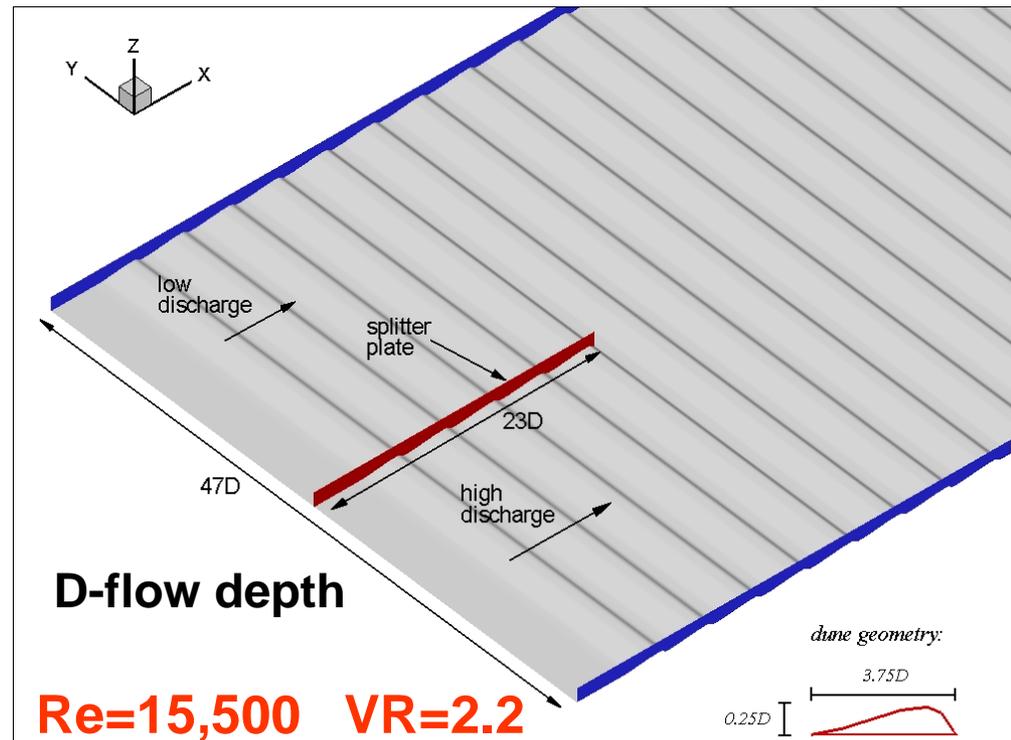
$$S(x) = \frac{\bar{c}_f \delta(x) U_c(x)}{2D(x) \Delta U(x)}$$

-characterizes stabilizing influence of bottom friction on ML development

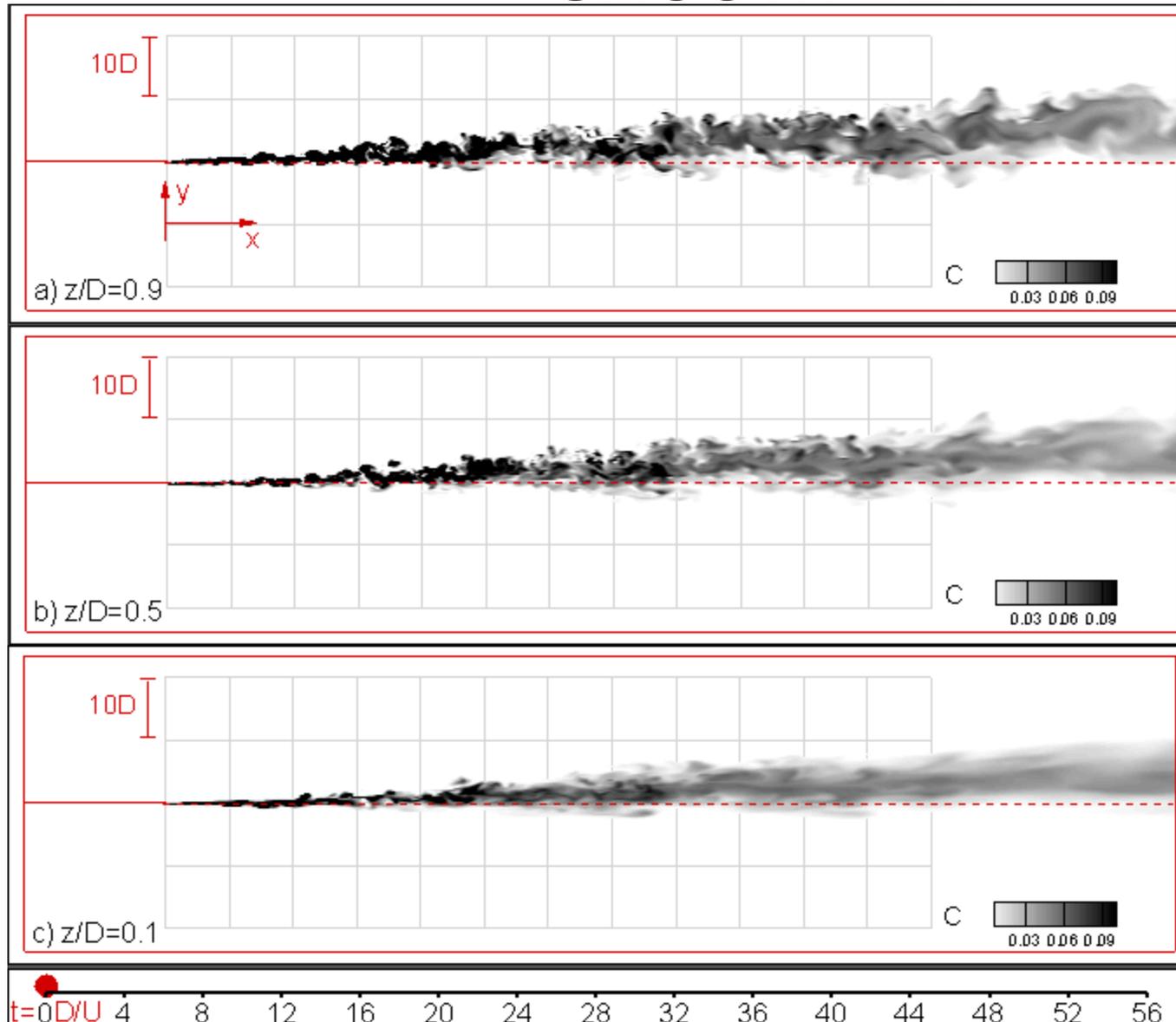
-**S=Sc** equilibrium

-Sc~0.1

-shift and width of ML
do not vary with x



Confluences between parallel streams SMOOTH BED

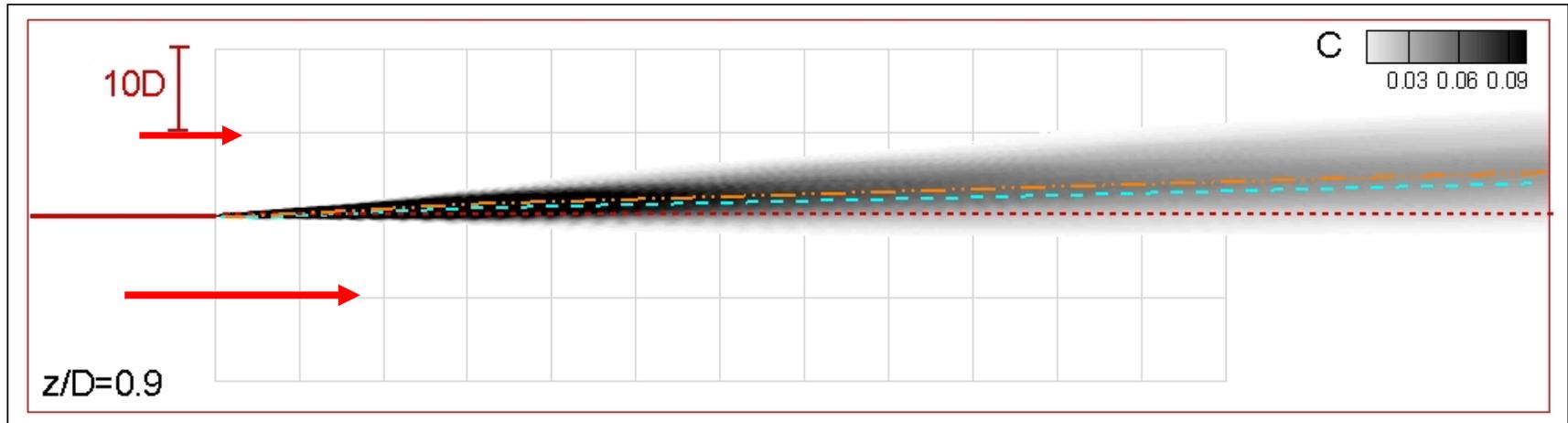


Max size=10-15D

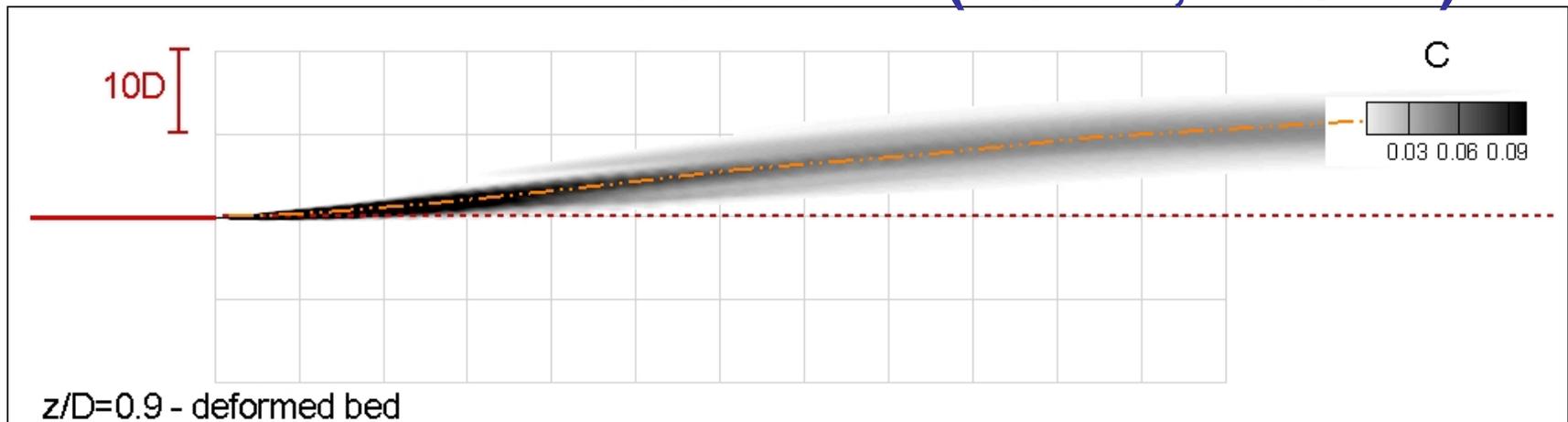
Confluences between **parallel** streams

Effect of bed friction

SMOOTH BED



ROUGH BED (DUNES, $H=0.25D$)



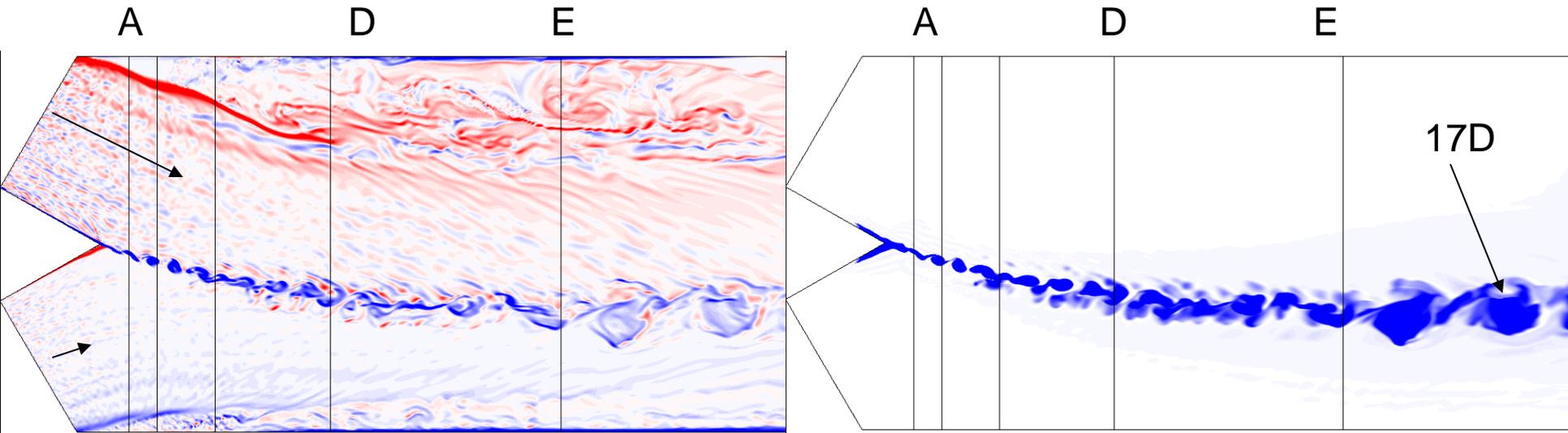
MIXING INTERFACES

-If the **angle** between the two streams is **large** and the **velocity difference** between the two streams is **high**, we get a more complex type of MI:

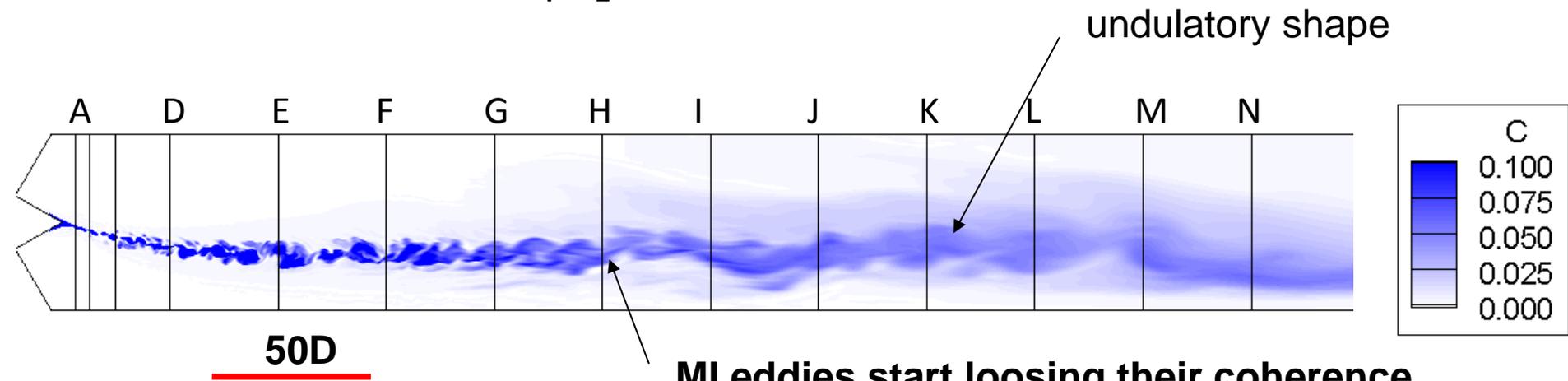
SHALLOW MIXING LAYER developing between two non-parallel streams

-Some interesting results obtained in an idealized geometry:

MI at a confluence between two non-parallel streams



$\alpha=60^\circ$ $D=\text{constant}$ $U_1/U_2=2$ $Re=200,000$



MI eddies start losing their coherence after shift of the MI becomes constant ($x \sim 250D$)

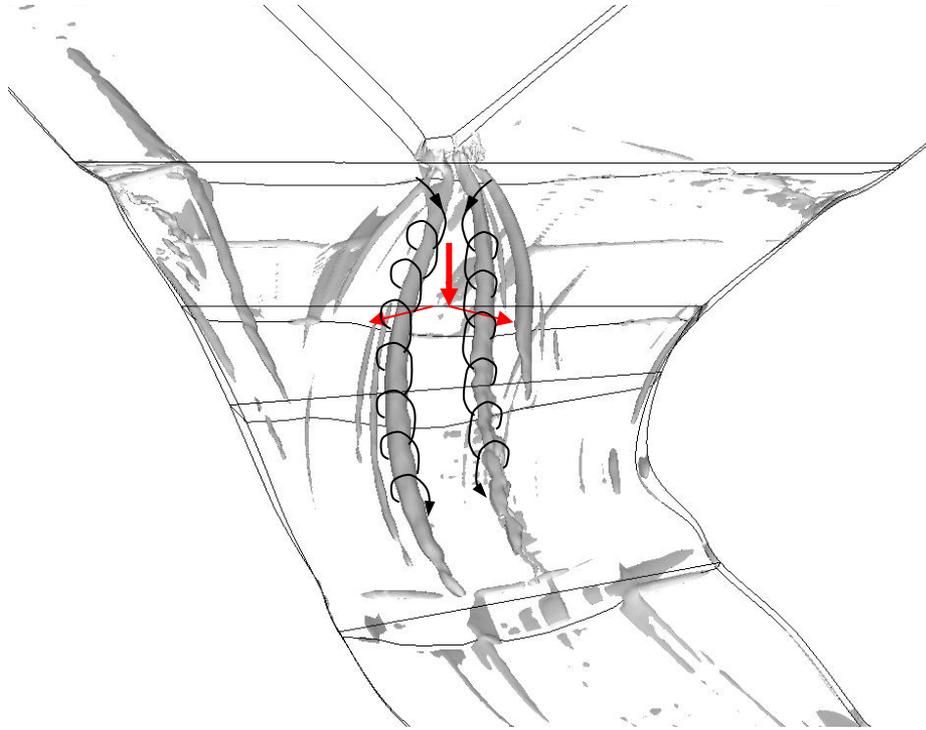
QUESTION

Besides MI eddies are there other types of large-scale coherent structures forming close to the MI?

-Generally, the answer is positive

-SOV cells can form in the vicinity of the MI, especially for confluences at which the degree of bed discordance is not very high

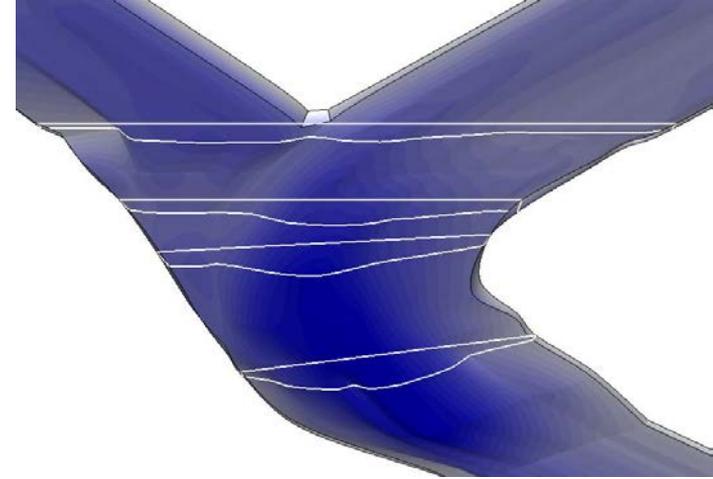
Mechanism responsible for formation of SOV cells



-Sediment entrained and transported by SOV cells is the main cause for the large dimensions of confluence scour holes ($D_{\text{scour}} \sim 5D$; Best & Ashworth, Nature 1997)

-Mass exchange processes and thermal mixing between the two streams are strongly affected by the SOV cells

MIXING INTERFACES

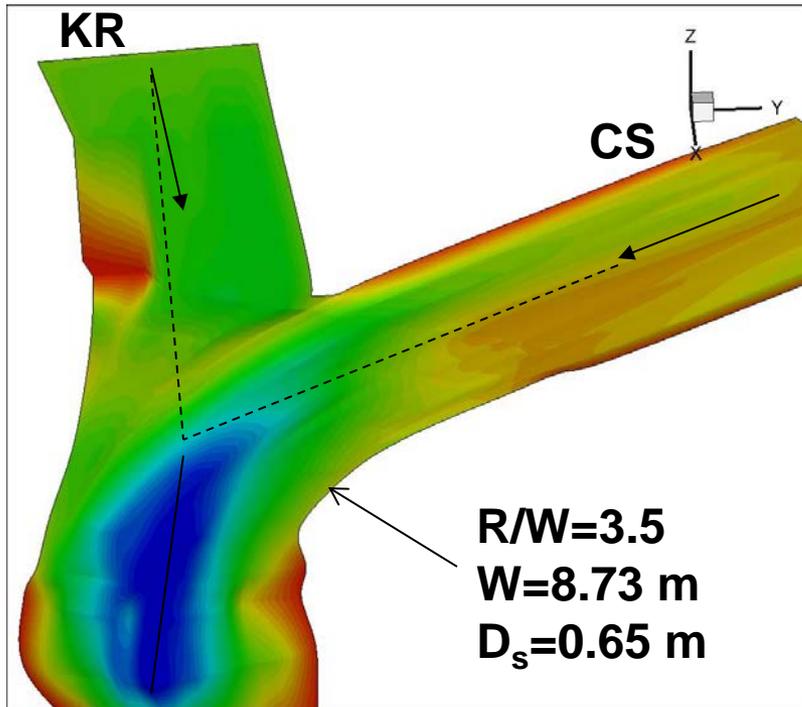


- Let us finally consider a **natural stream confluence**
- Small river confluence in Illinois
- Asymmetrical confluence with **concordant bed**
- Angles: 0° , 60°
- Field data available for validation (Rhoads & Sukhodolov, 2001, 2004, 2008)

CONFLUENCE BATHYMETRY

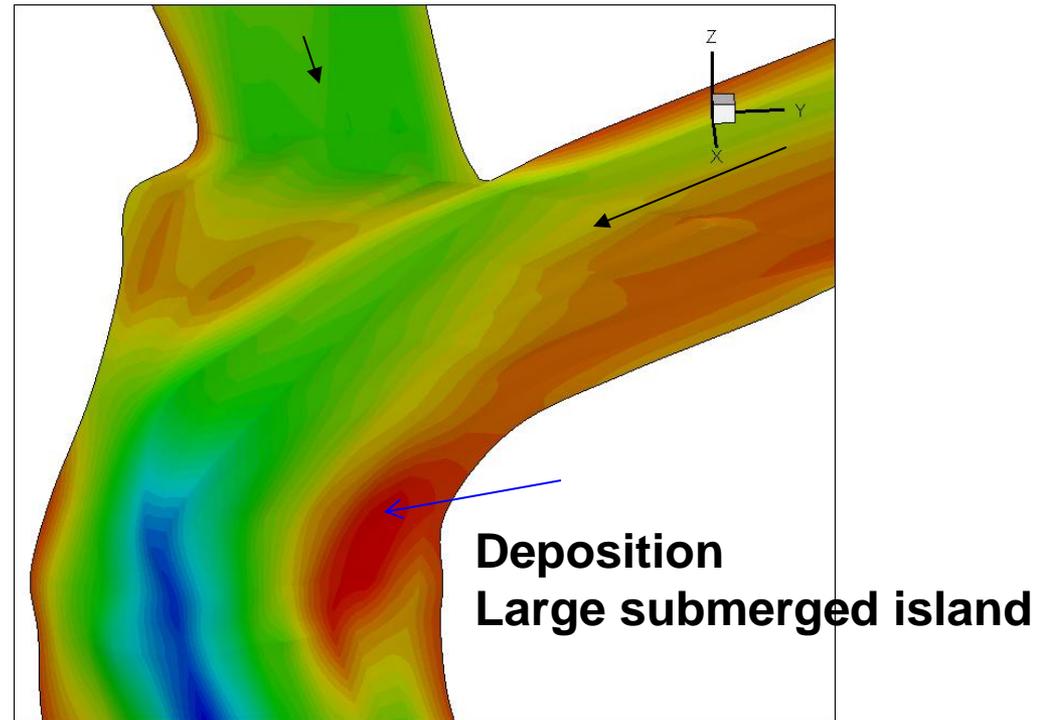
CASE 1 ($V_r \sim 1$)

$Re \sim 166,000$ ($D=0.36\text{m}$ $U=0.45$ m/s)



CASE 2 ($V_r=5.5$) - 1 year later

$Re \sim 77,000$ ($D=0.23\text{m}$ $U=0.34$ m/s)

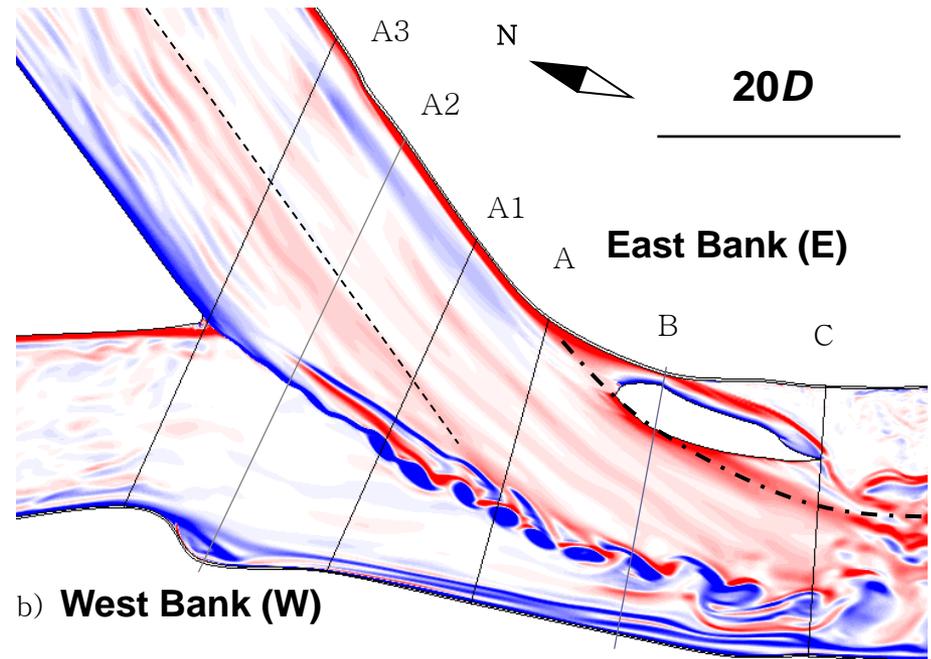
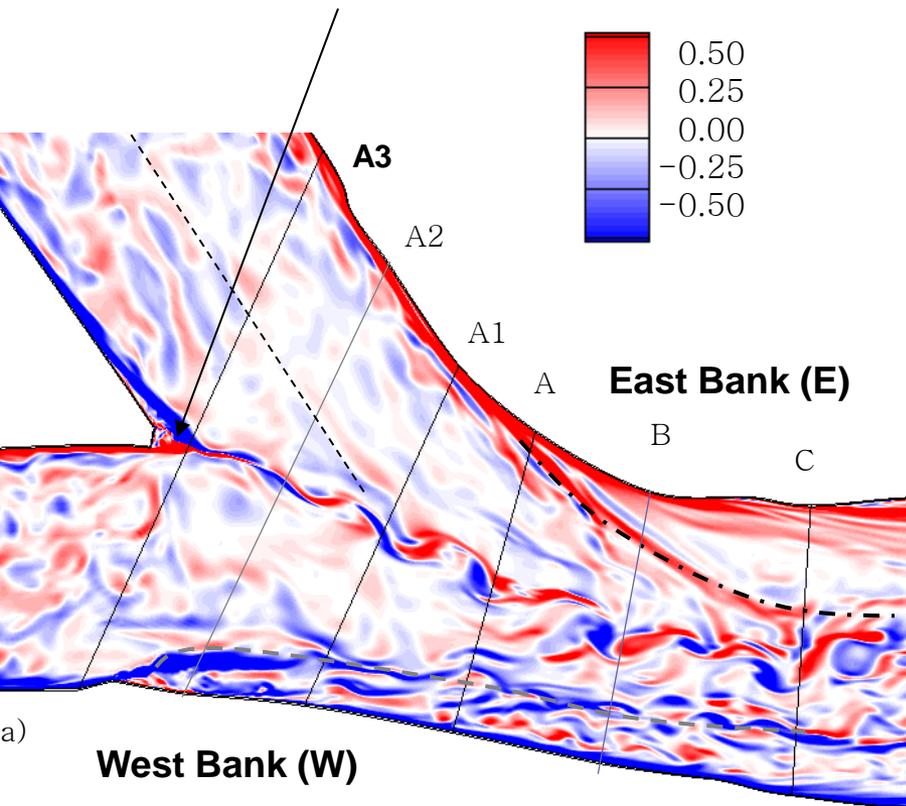


Vortical structure close to free surface

VR~1

VR=5.5

Wake region



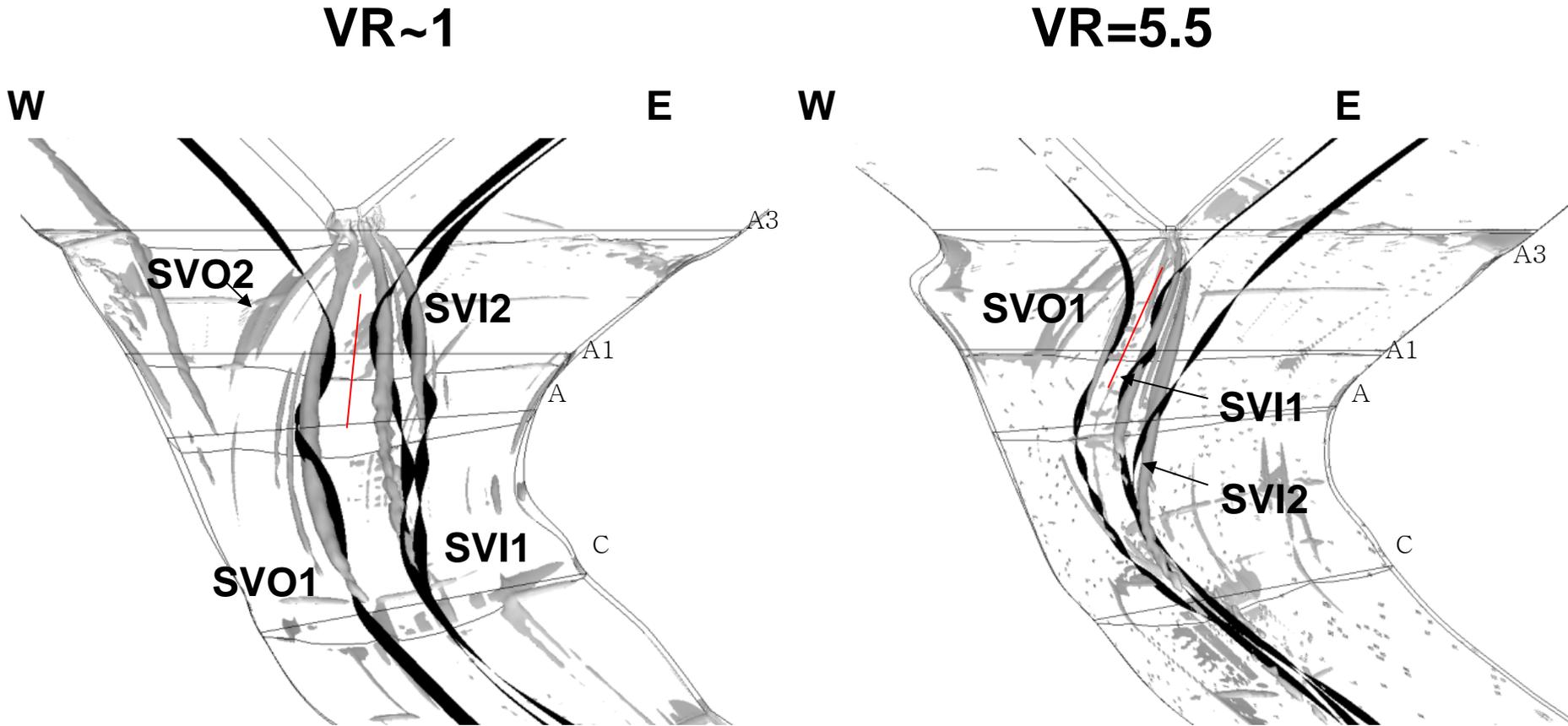
Case 1: MI is in **wake mode**
(von Karman vortex street)

Case 2: MI is in **KH mode**

-MI contains eddies with opposite
sense of rotation shed from wake region

-MI contains co-rotating eddies

Large-scale eddies below the free surface MEAN FLOW

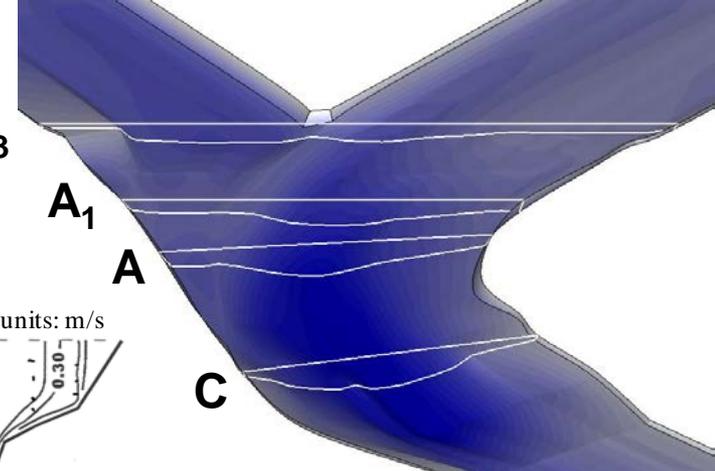


-Primary SOV cells are counter-rotating

-VR=5.5: SOV cells are much more coherent on the high momentum side

VALIDATION (VR~1)

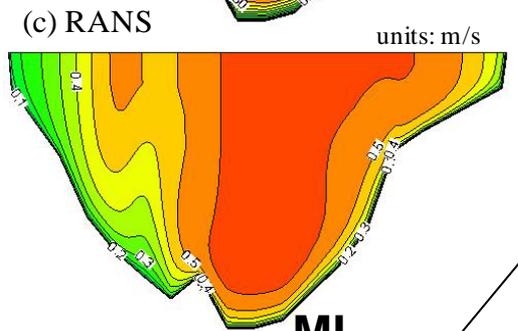
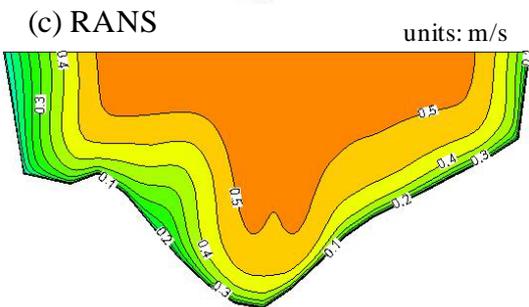
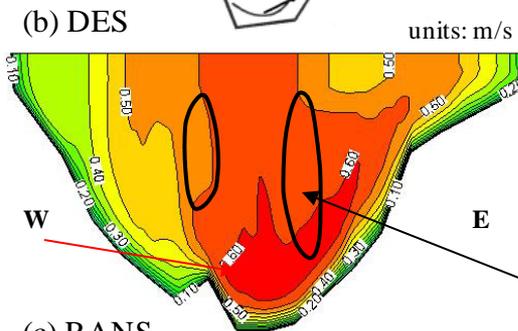
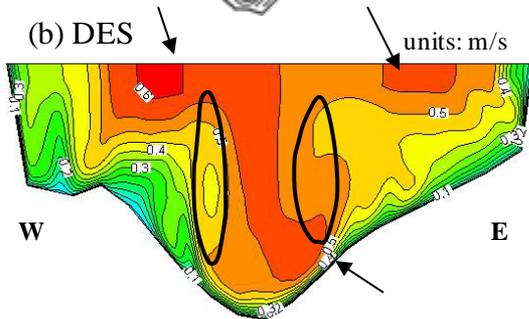
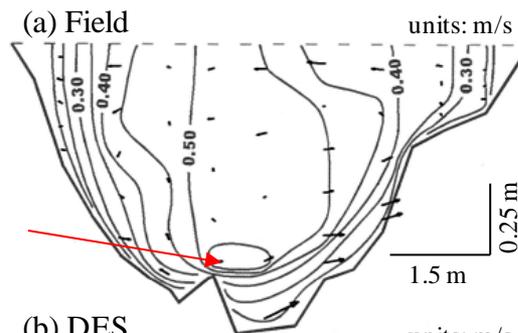
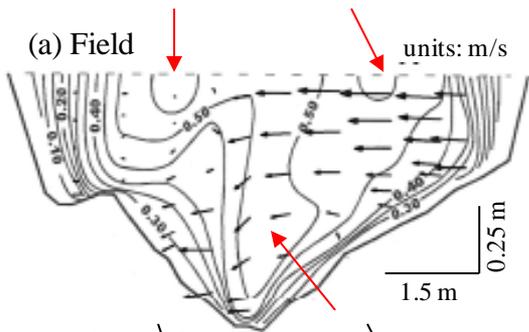
Streamwise velocity



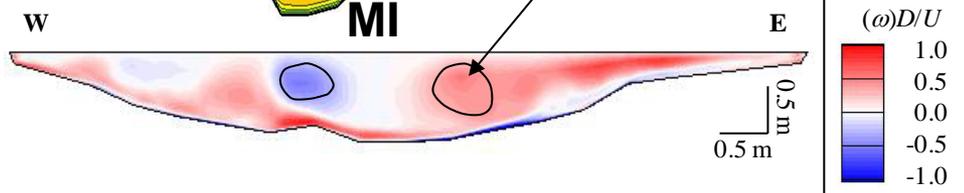
Section A

Section C

Scale 1:6

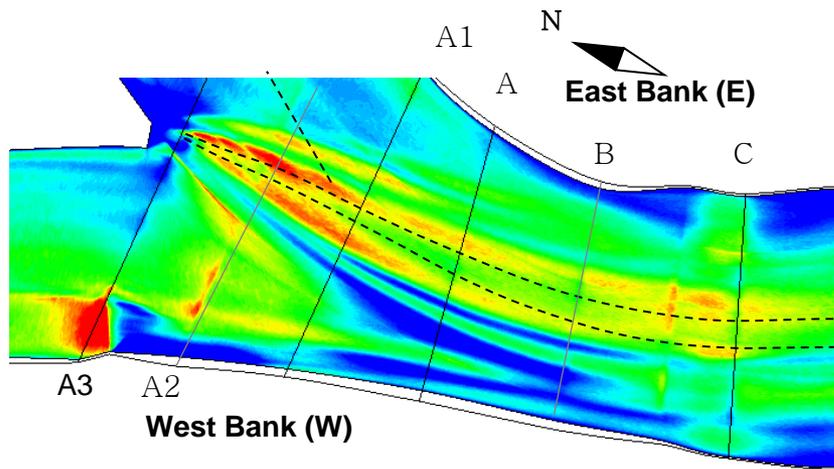


Out-of-plane vorticity
Section C – scale 1:1

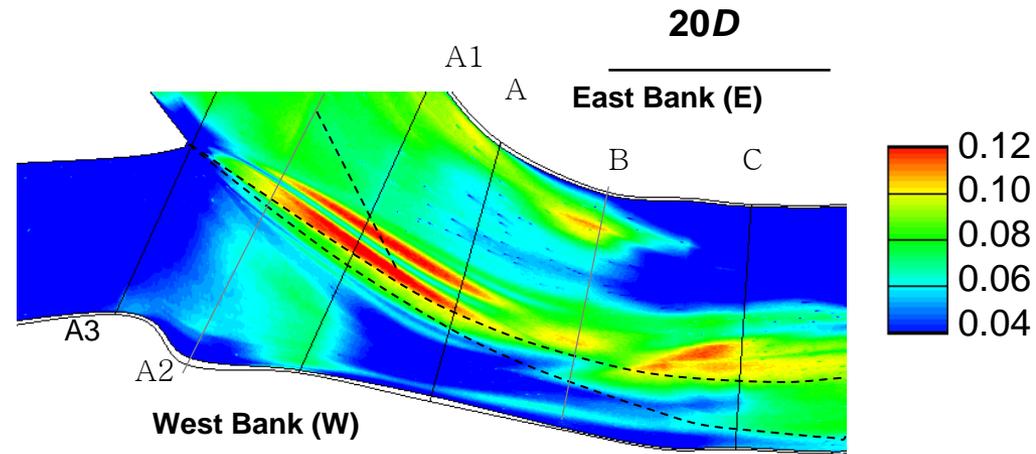


SOV cells are
much stronger
in DES

Bed friction velocity: Mean flow

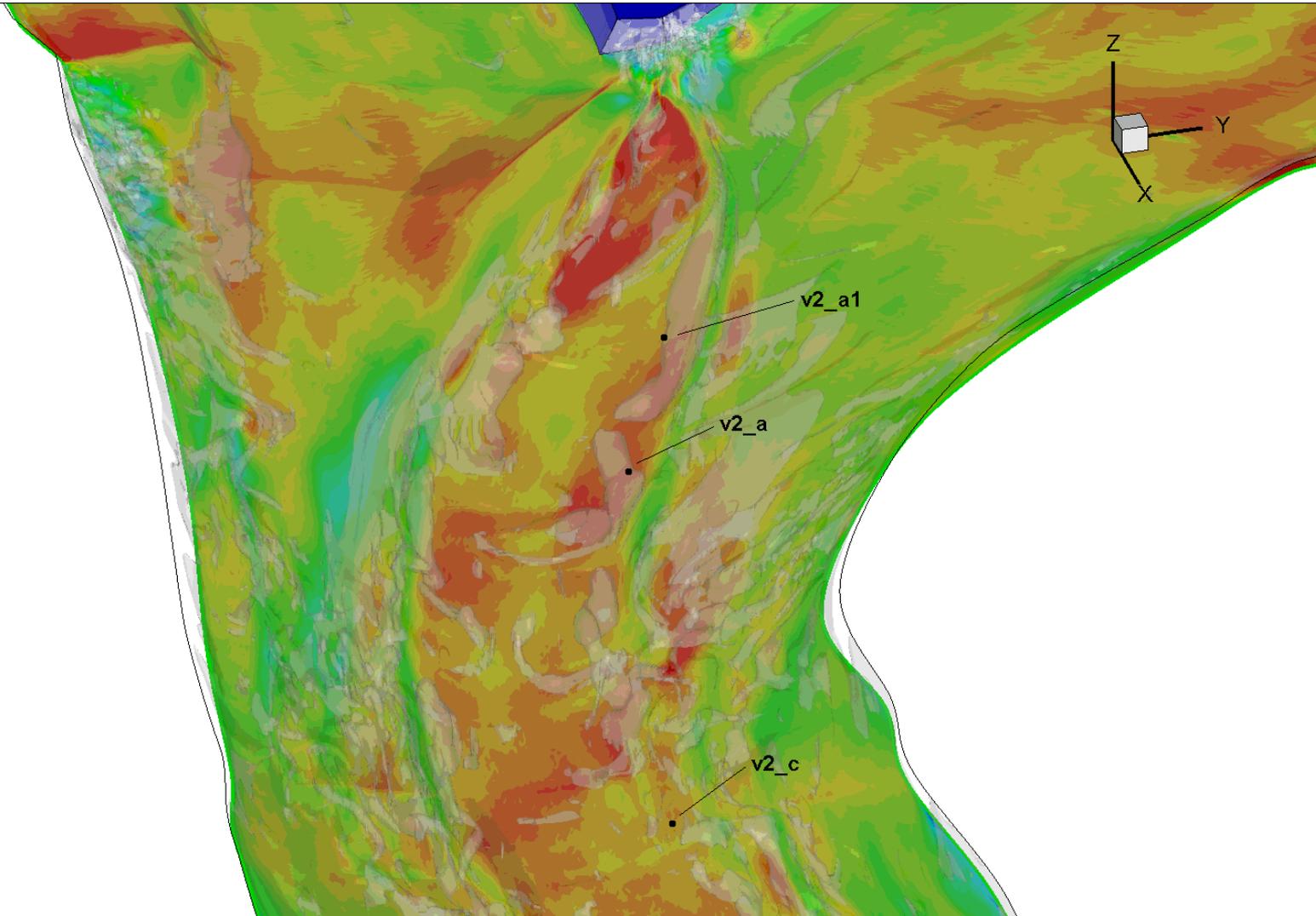


VR~1



VR=5.5

Bed friction velocity: Instantaneous flow



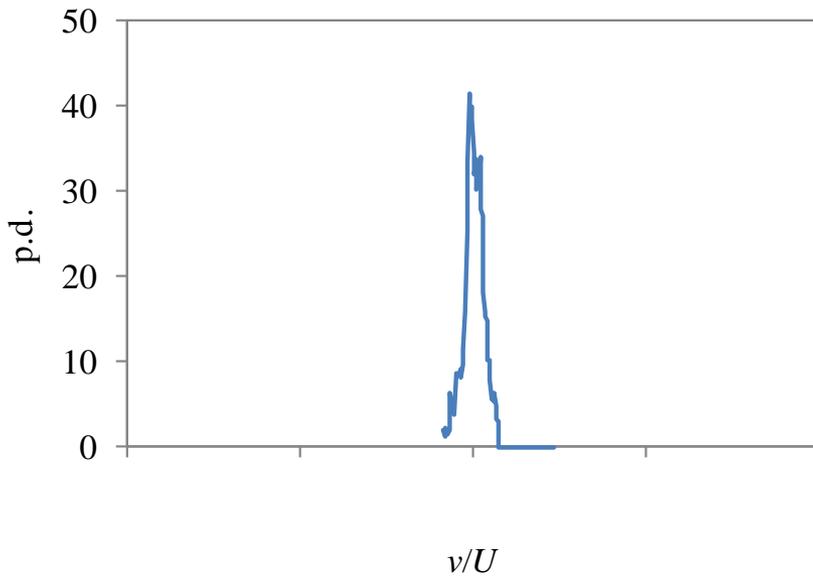
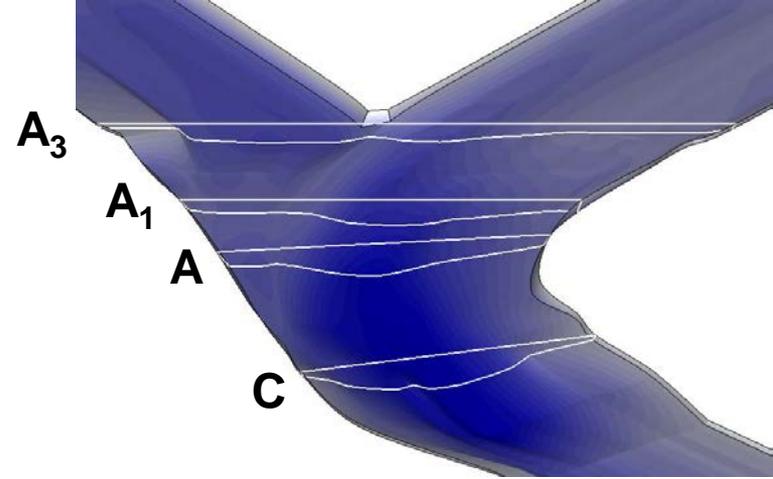
VR~1

Why do some of the SOV cells induce large bed friction velocities in the region where the two streams collide?

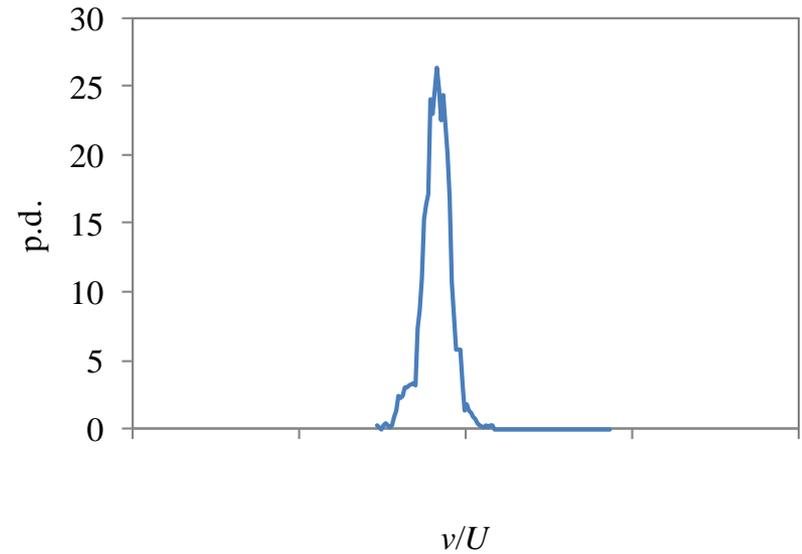
- Incoming flow has to lose rapidly a large amount of transverse momentum as it approaches the MI
- Strong adverse pressure gradients are created as the cores of high streamwise velocities in the two streams approach the sides of the MI
- This situation is similar to **junction flows** in which the necklace vortices form in a region of strong flow deceleration
- The necklace vortices have a large capacity to entrain sediment because they are subject to **large-scale bimodal oscillations**

Velocity histograms outside the main SOV cells:

CASE 1 (VR~1)



Incoming stream (CS)



Mixing Interface

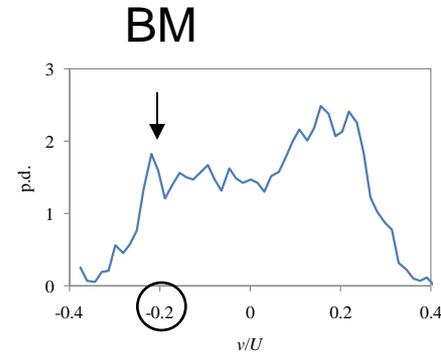
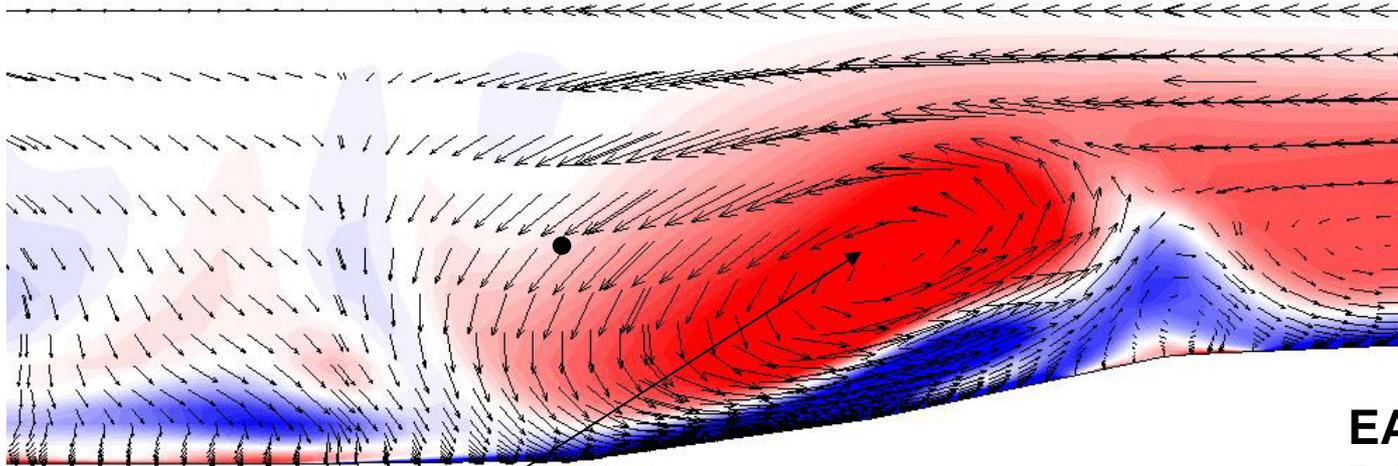
Velocity histograms contain only one peak

Velocity histograms inside SVI1: $VR \sim 1$

Section A_1

Bank mode

Histograms of v



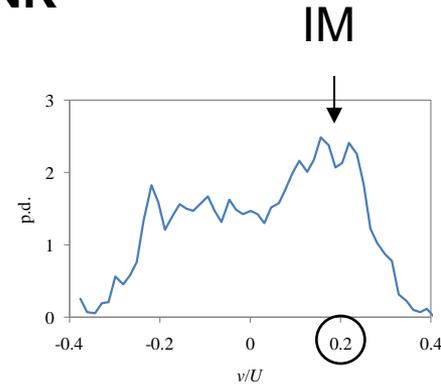
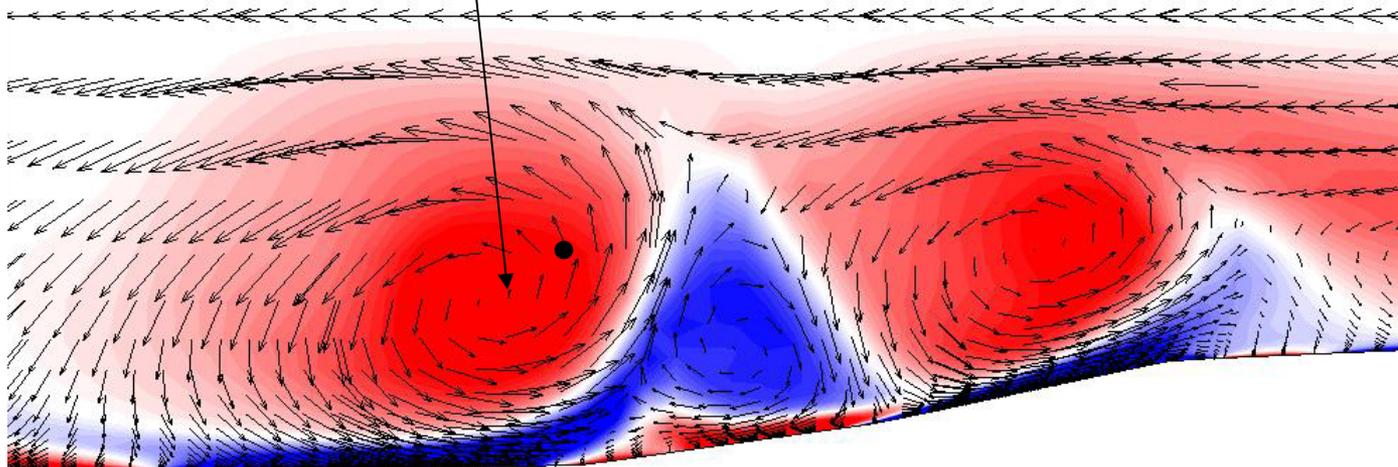
EAST BANK

MI

SVI1

Interface mode

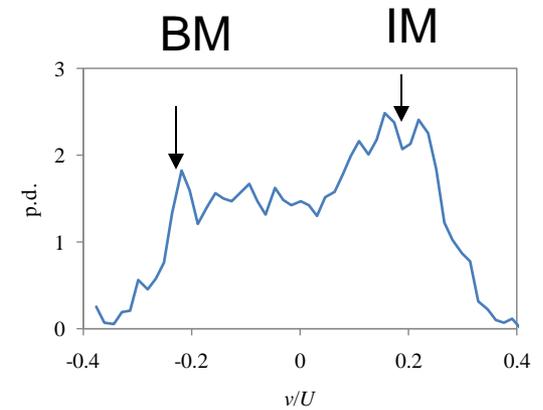
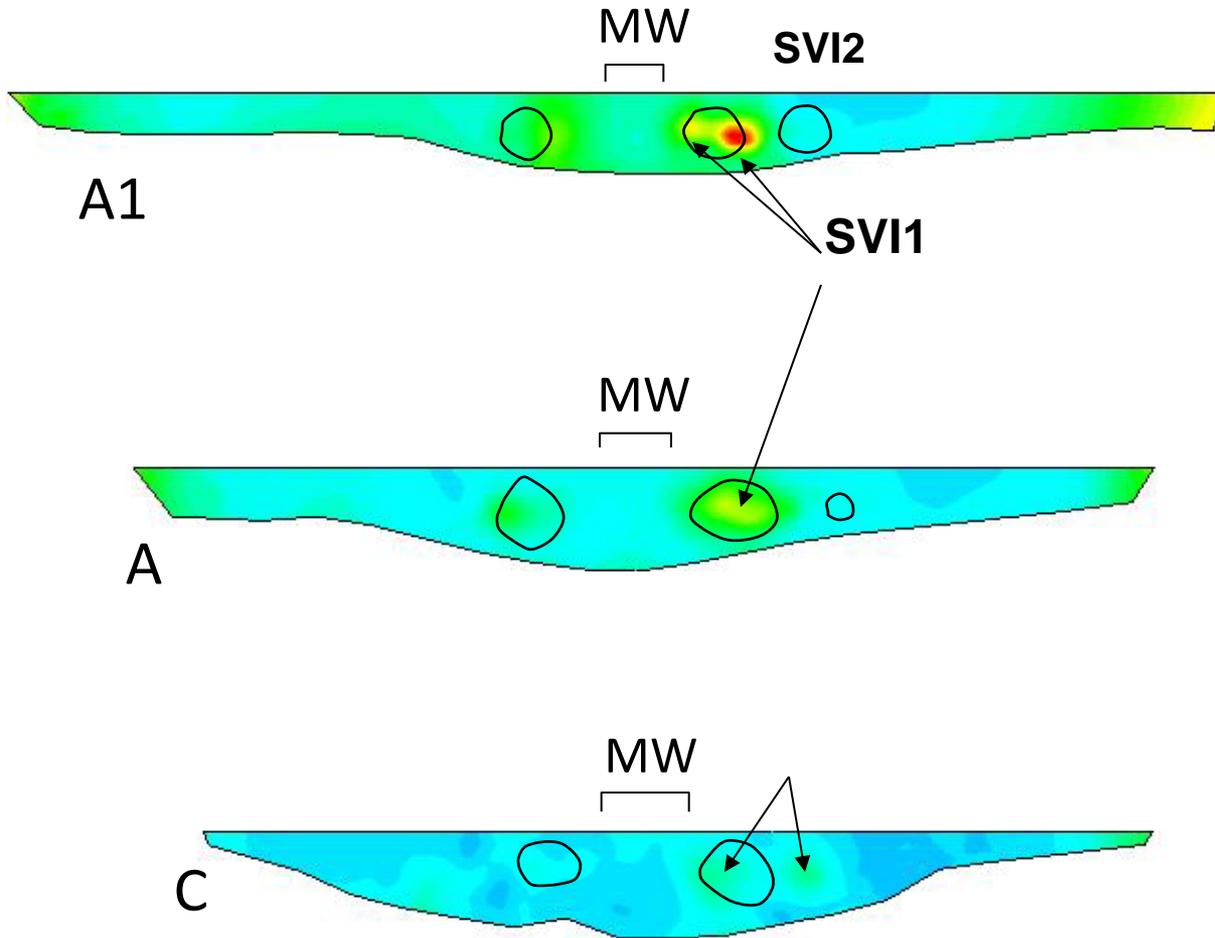
IM



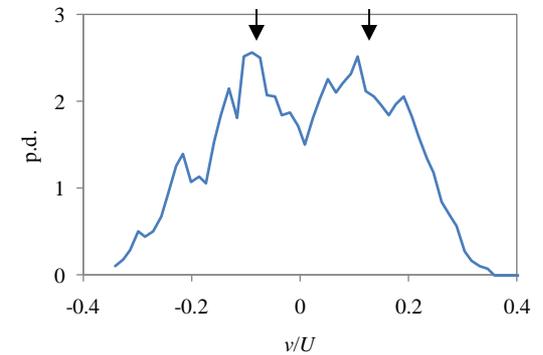
SVI1 is subject to bimodal oscillations

Streamwise variation of intensity of bimodal oscillations:

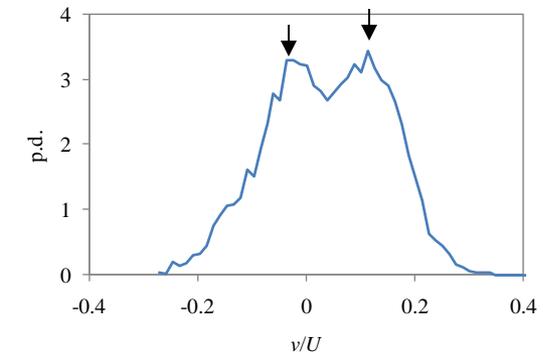
Pressure rms fluctuations



A1



A



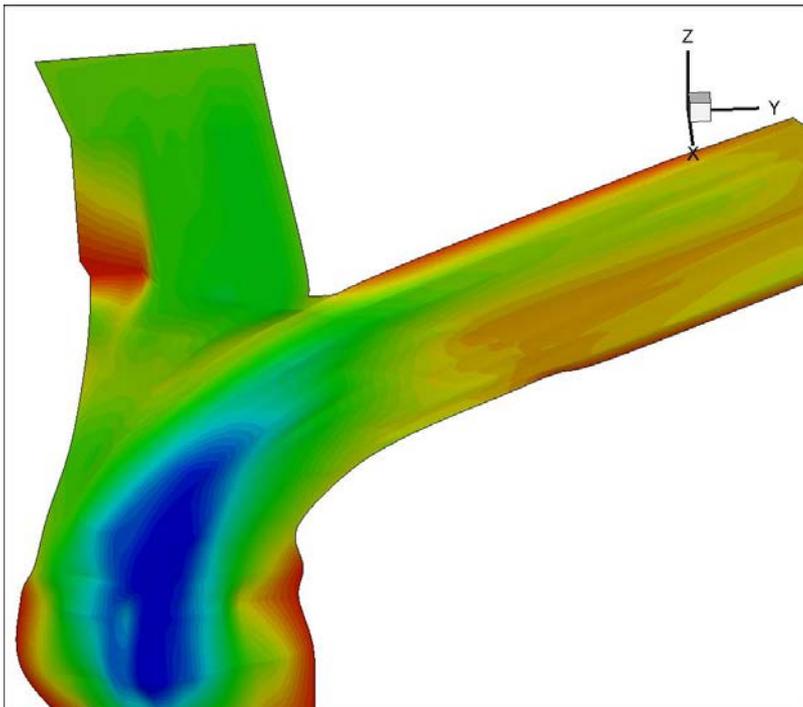
C

Eventually, transition to 1-peak shape will occur in the downstream channel

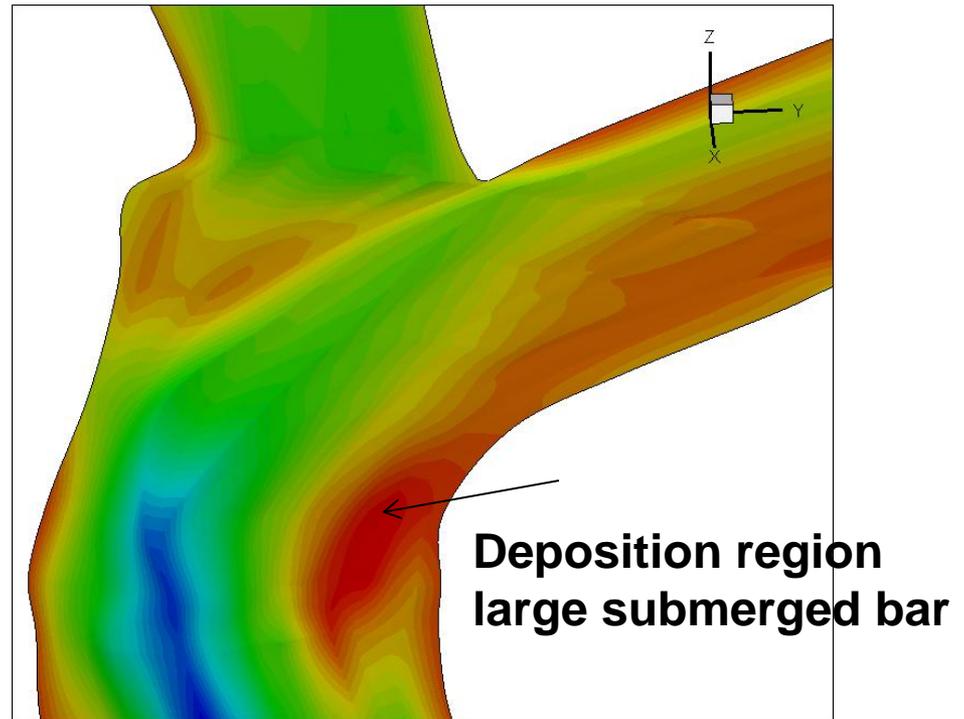
Do SOV cells play a role in formation of a large deposition bar at the inner bank?

BATHYMETRY

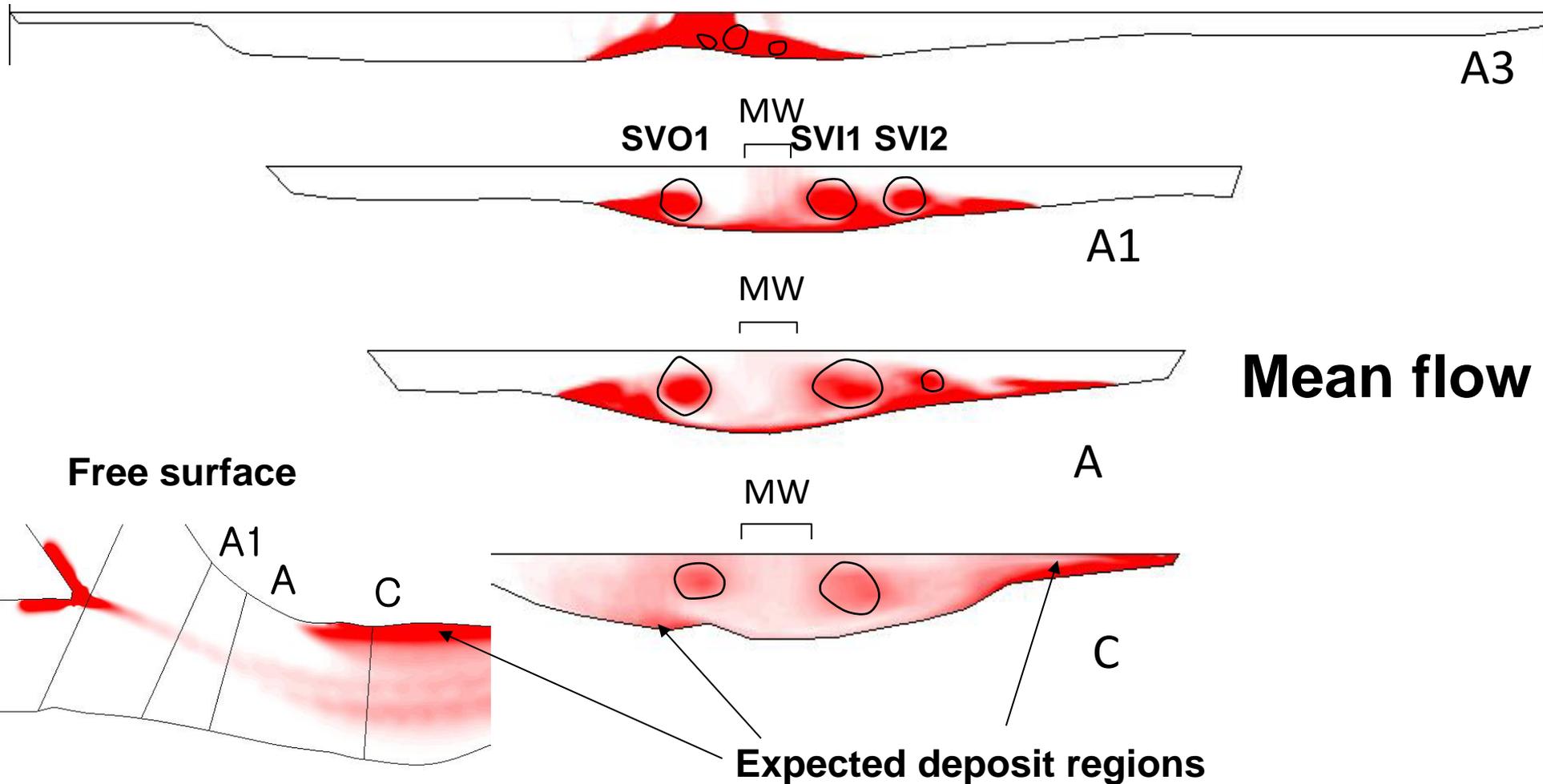
CASE 1



CASE 2 (1 year later)



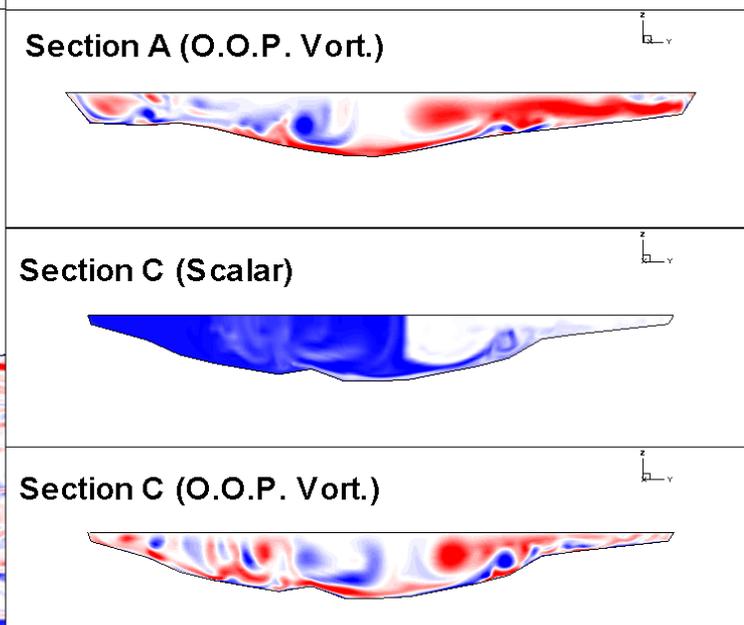
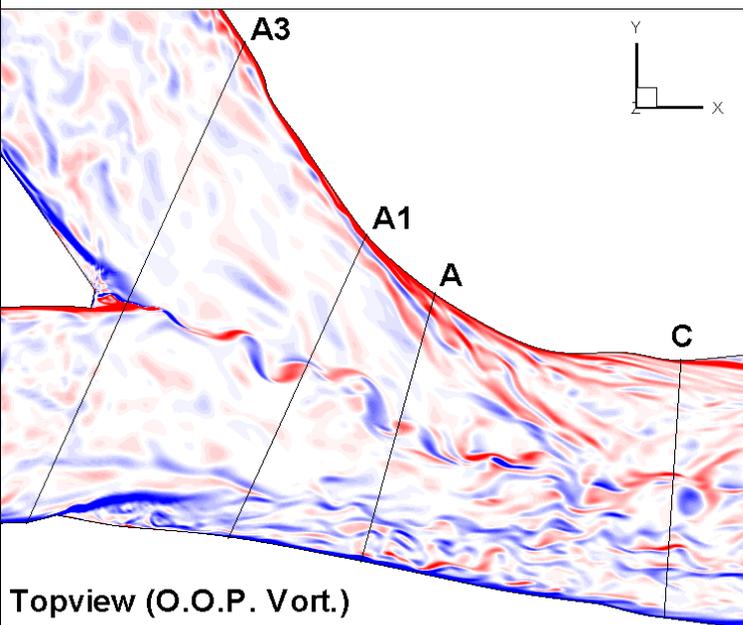
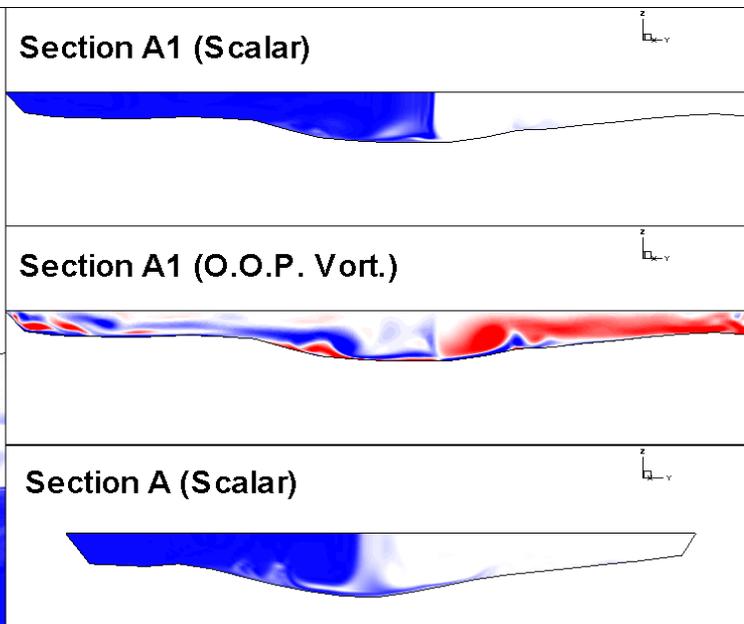
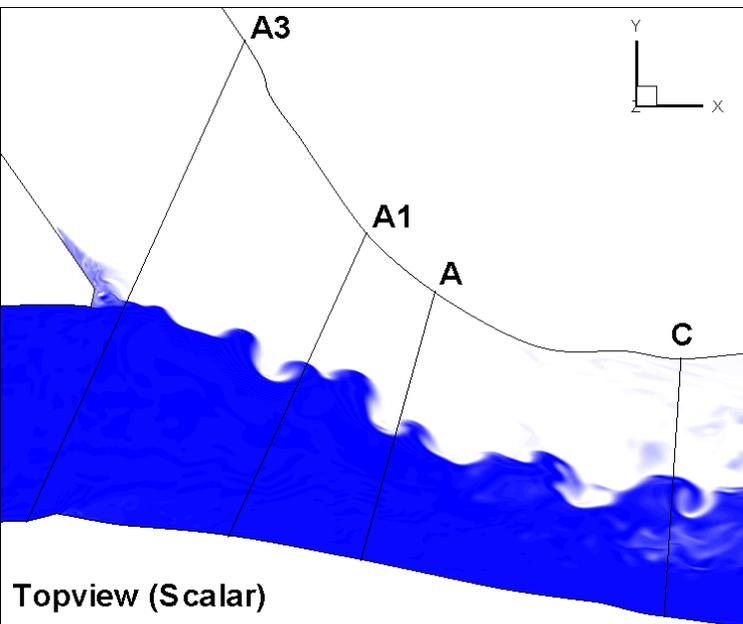
TRACKING OF PASSIVE SCALAR INTRODUCED AT THE UPSTREAM JUNCTION CORNER: $VR \sim 1$



SOV cells act as a pump of momentum and mass which besides entraining sediment from beneath, extract fluid and suspended sediment from the MI and advect them downstream of the CHZ

Can SOV cells affect mixing within the confluence?

THERMAL MIXING: $VR \sim 1$



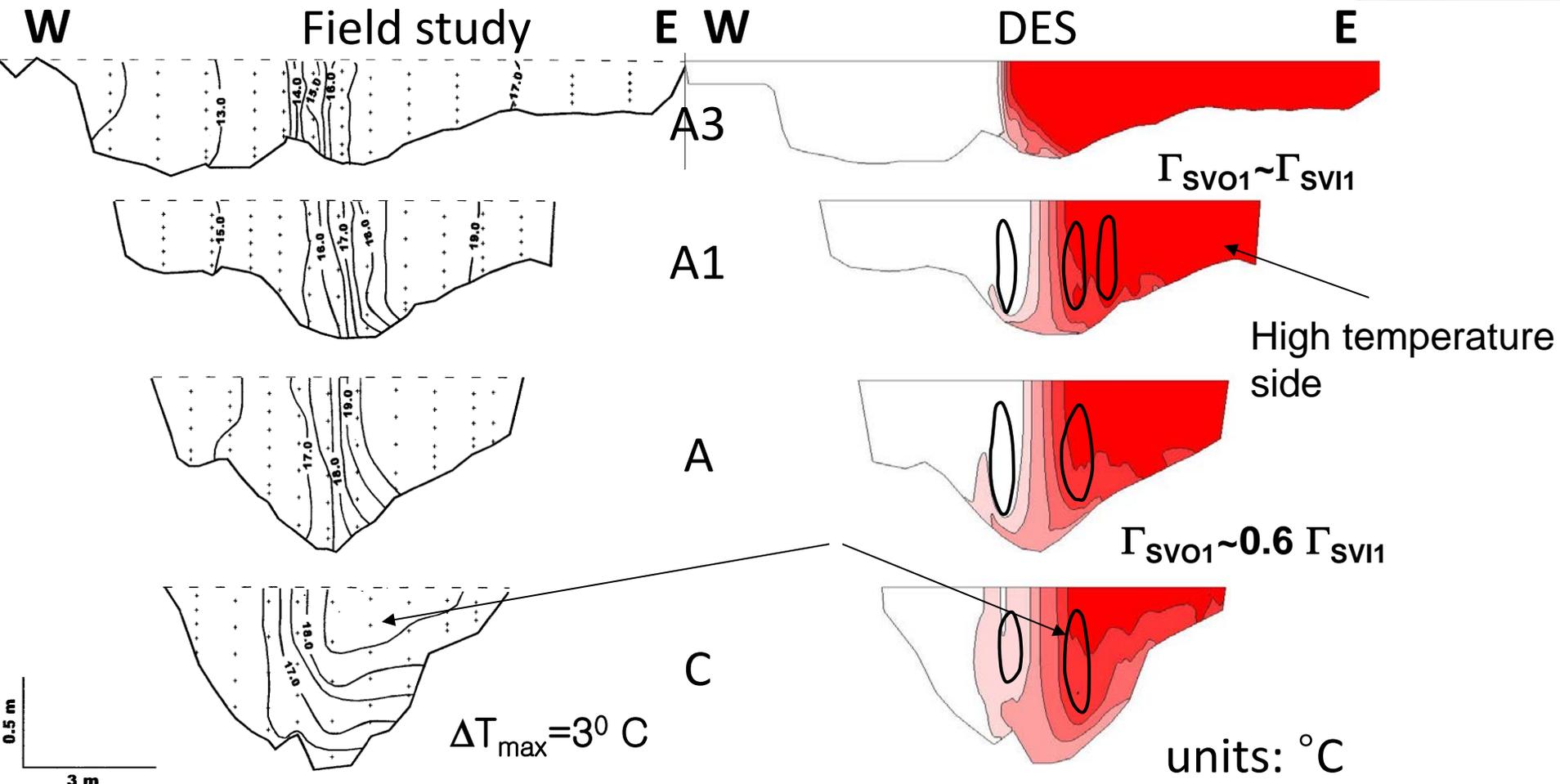
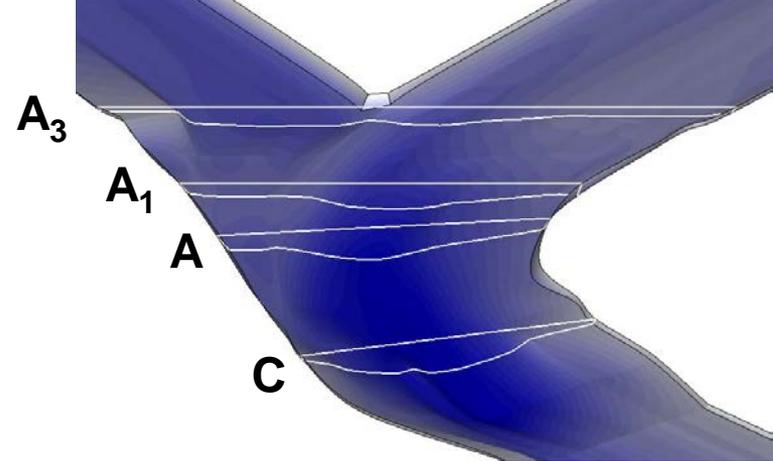
Large-scale engulfing by the ML eddies

Strong near-bed intrusions of low temperature fluid on East side driven by bimodal oscillations of SVI1

BLUE: low temp West side

VALIDATION: $VR \sim 1$

TEMPERATURE FIELD

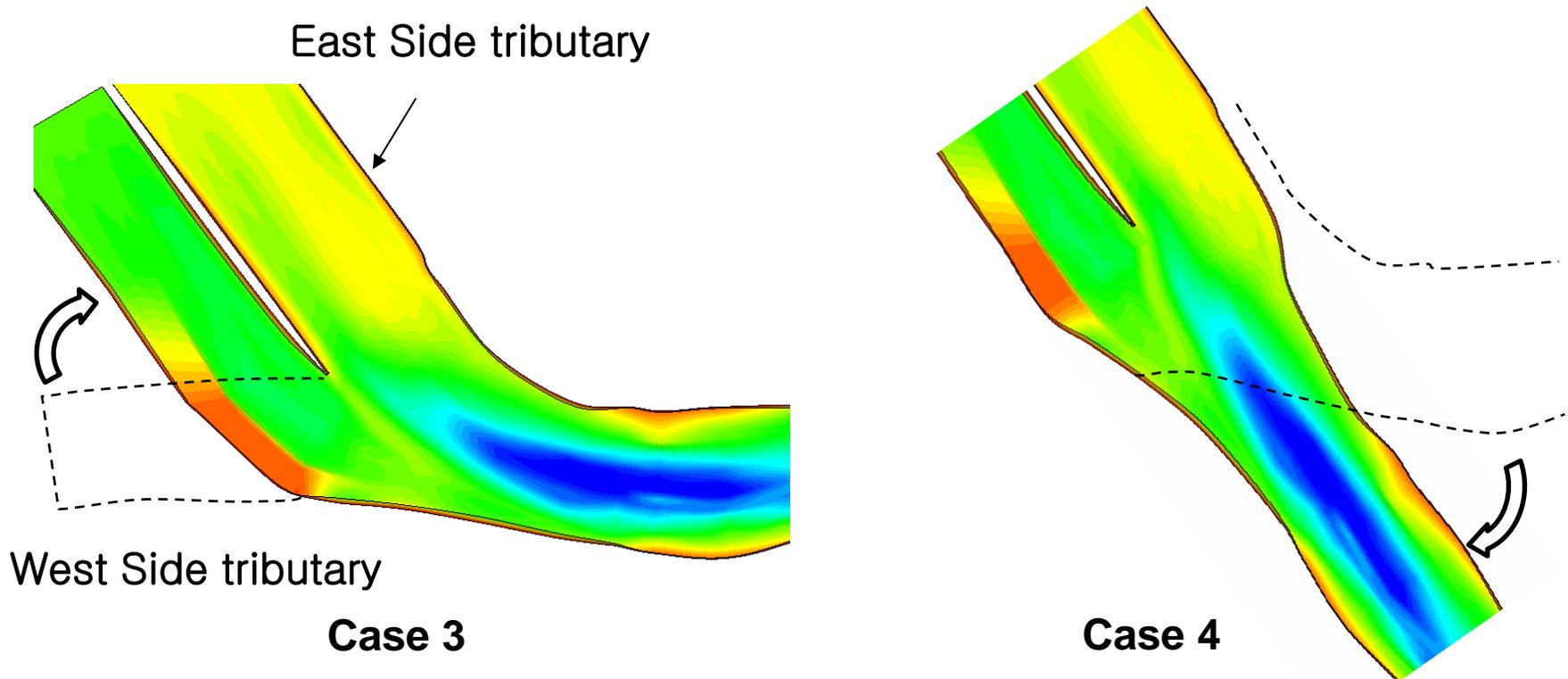


Conditions needed for SOV cells to form

-large angle between the two streams?

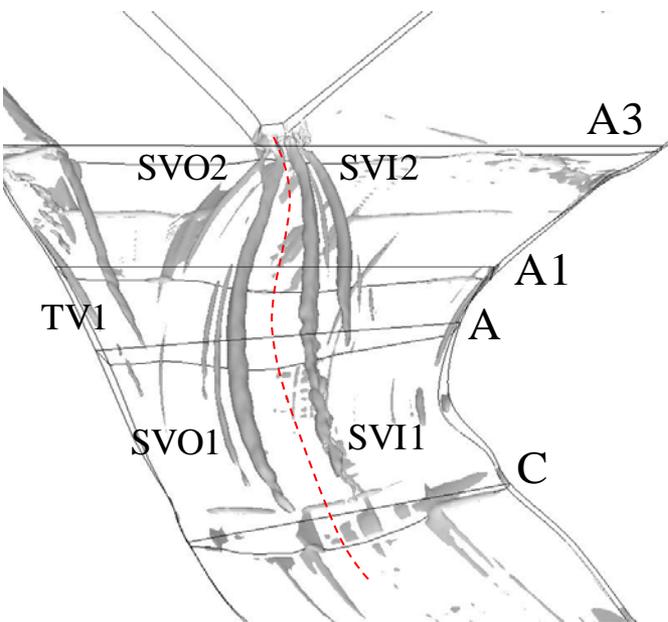
-large angle between at least one of the incoming streams and the downstream channel?

We will conduct a numerical experiment



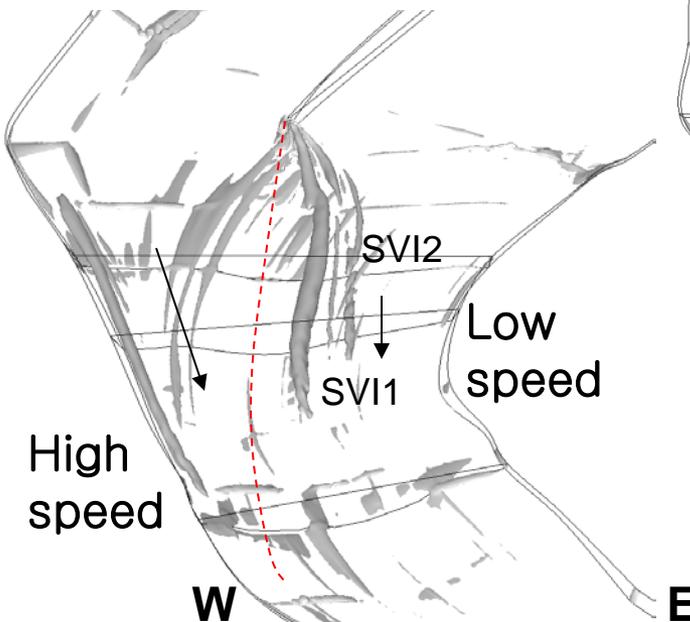
Streamwise Oriented Vortical Cells

Case 1



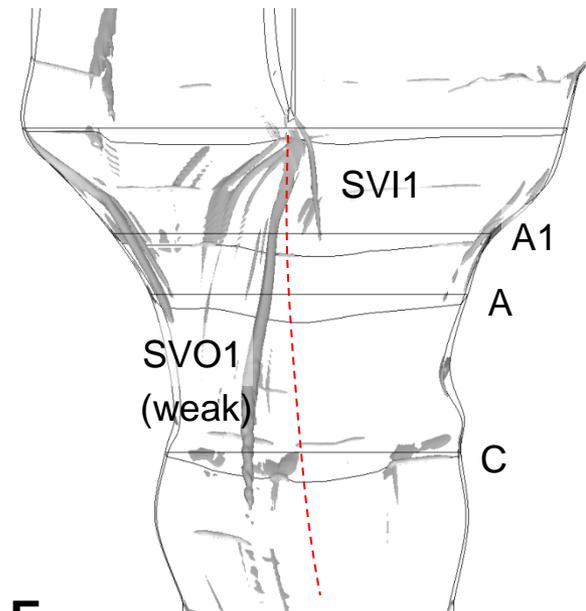
MI: wake mode

Case 3



MI: KL mode in ds part

Case 4



MI: wake mode

- weak MI eddies
- weak SOV present

- left tributary pushes into MI (SVI1)
- right tributary pushes into bank

Why sometimes the two streams do not appear to mix for large distances from the confluence apex?

A famous example (Amazon): Confluence of Solimoes and Negro Rivers

-Solimoes River contains sediment eroded from the Andes Mountains

-Negro River contains low sediment but high organic matters from the forest (black tea color)

-the two streams are also characterized by significant differences in temperature (4°C), nutrient and oxygen content

-the two rivers run side by side for 6 Km before they start mixing!



Confluence of Solimoes and Negro Rivers

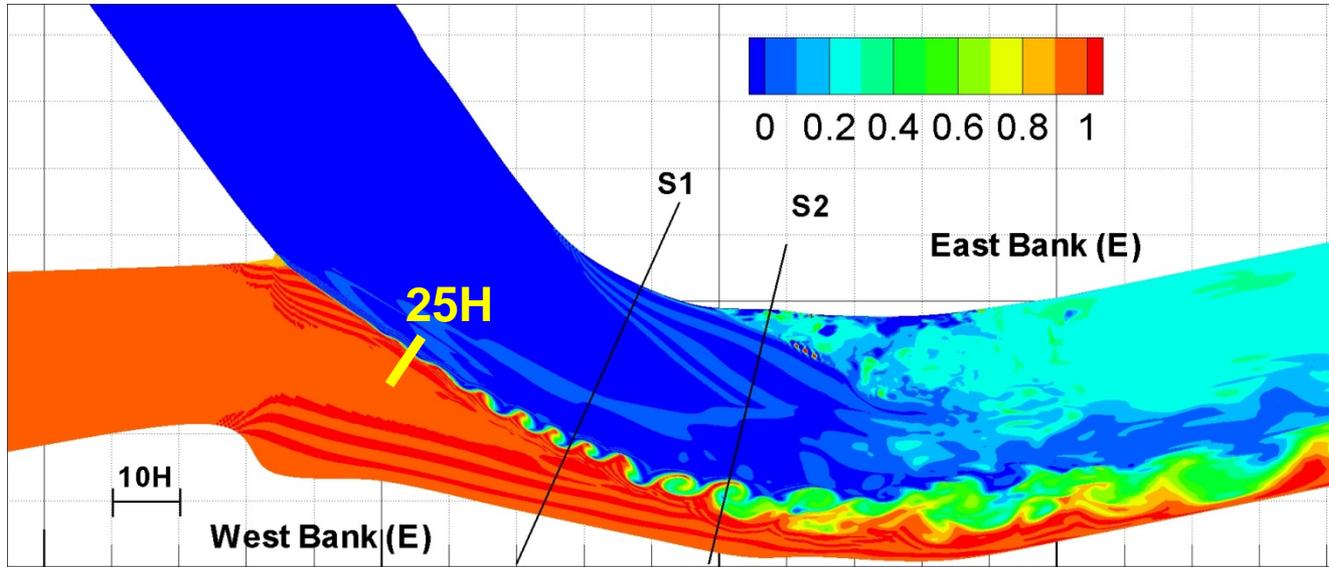


-Differences in temperature and suspended sediment concentration means that **stratification effects may be important !**

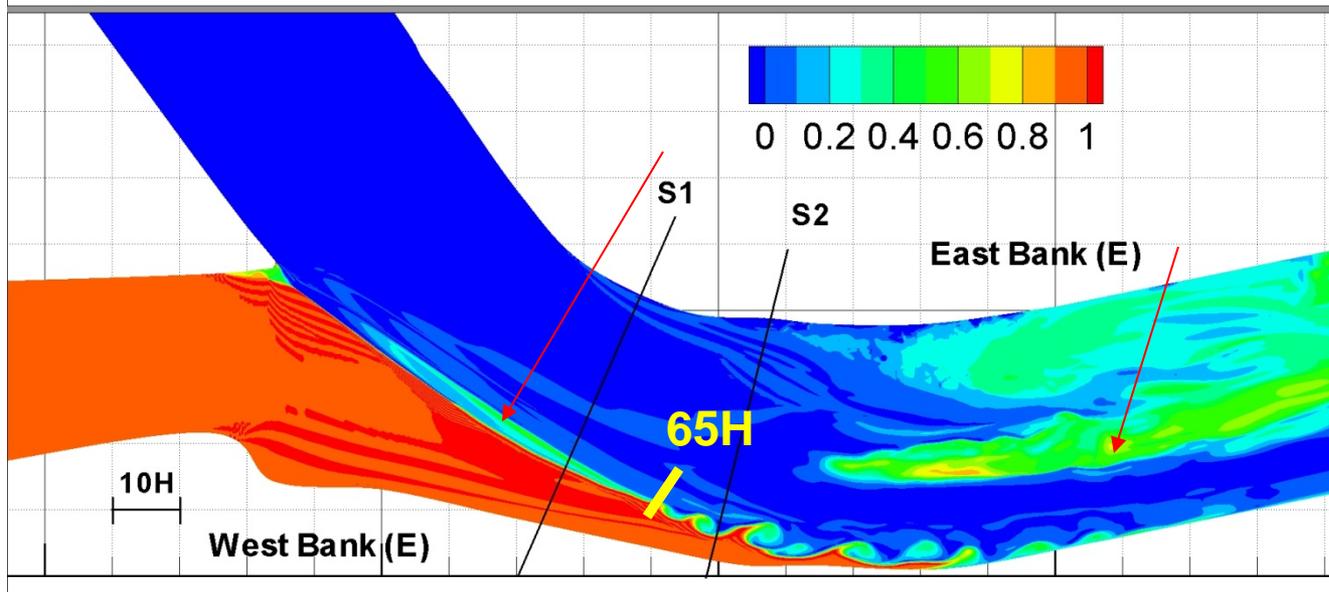
-Can stratification delay formation of large-scale eddies within MI?

-Back to our small confluence in Illinois

STRATIFICATION EFFECTS (Case 2, $V_r=5.5$)



$Ri=0.0$



$Ri=0.1$

Confluences with a strong discordance in bed levels and complex bathymetry features

-Are we still talking about a MI that is just a more complex case of a shallow mixing layer?

-Do SOV cells still form?

Confluence of Ebro and Segre Rivers

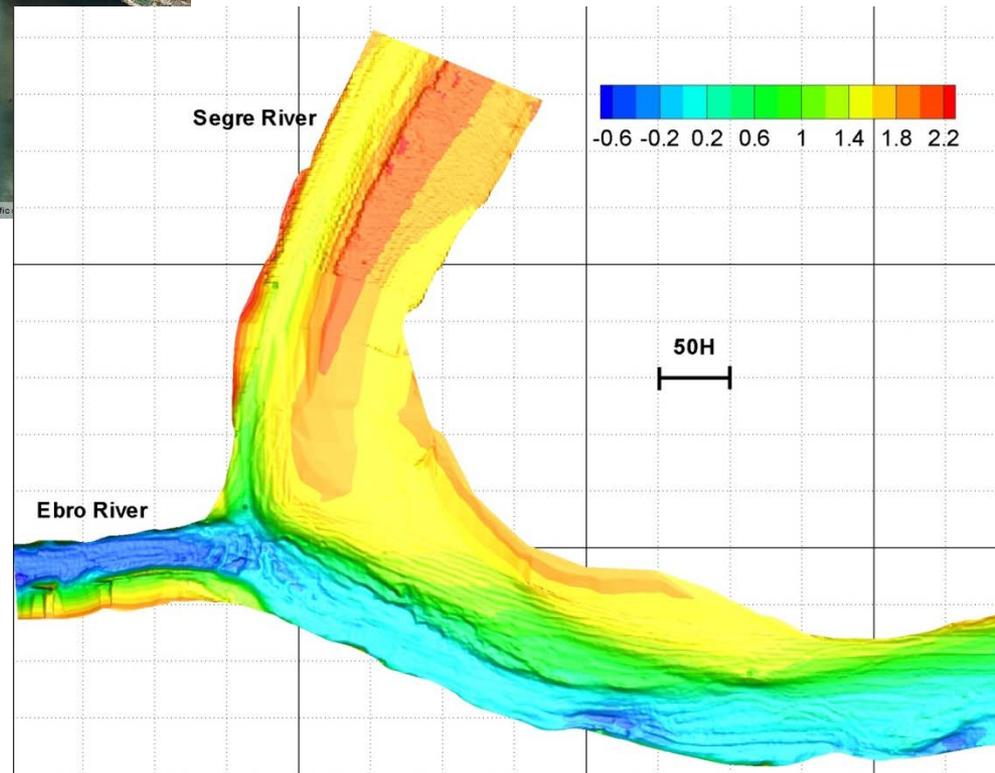


$$Q_{\text{EBRO}} \sim 400\text{-}500 \text{ m}^3/\text{s}$$

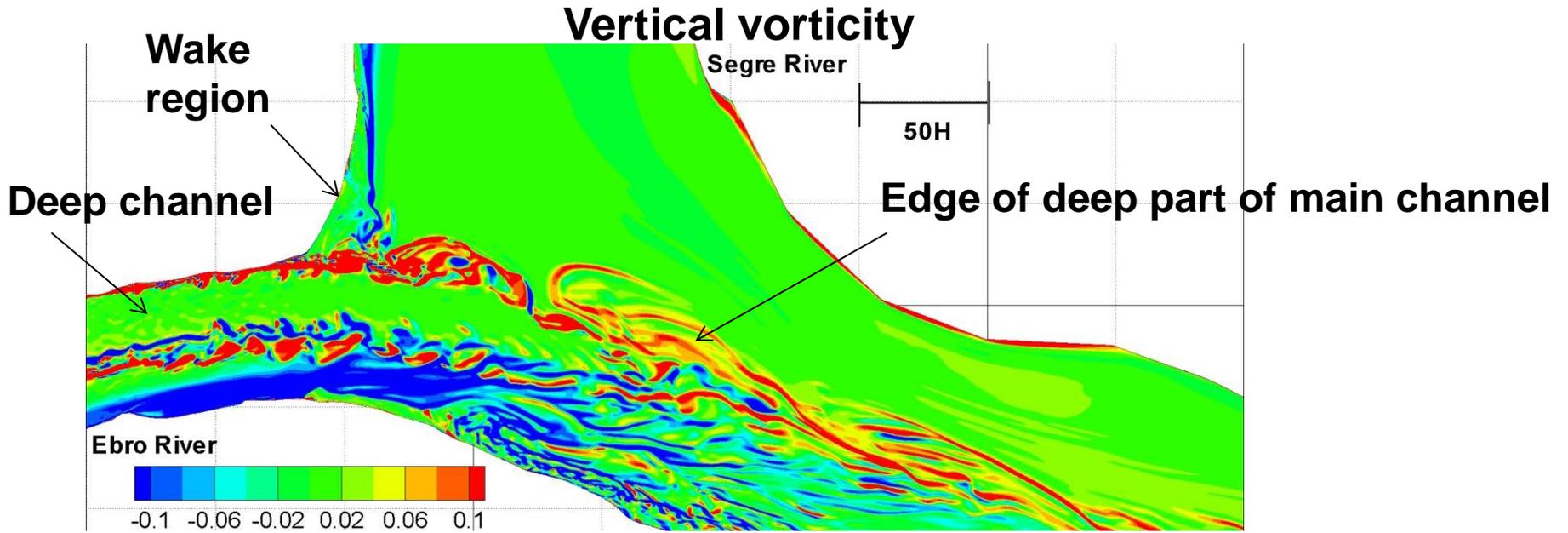
$$Q_{\text{SEGRE}} \sim 15\text{-}90 \text{ m}^3/\text{s}$$

$$QR = 2.5 \text{ to } 13 \text{ !!!}$$

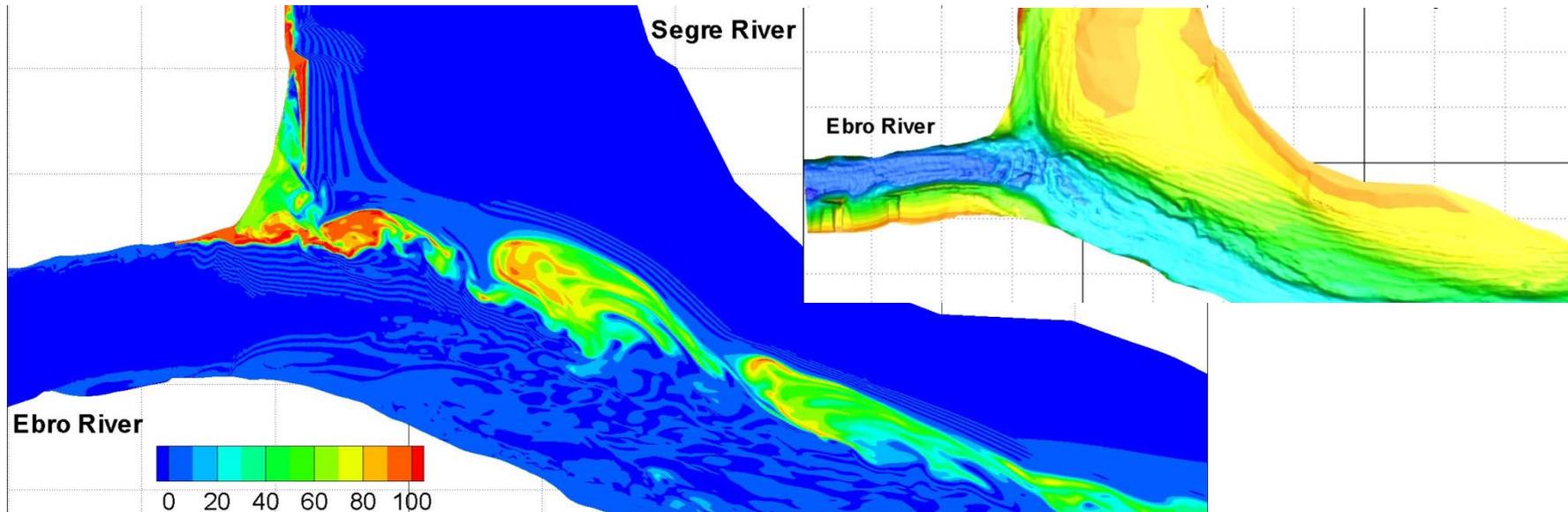
Bathymetry data courtesy of
J. Dolz and J. Prats-Rodriguez



Confluence of Ebro and Segre Rivers $QR=2.5$



Passive scalar introduced at confluence apex



FINAL REMARKS

-Eddy resolving techniques can be used to better understand flow physics and to test hypotheses related to mixing and transport in natural streams

Challenges:

-Simulate larger-scale systems

-Bed-morphology changes

-Integrate ecological modeling